



Large daytime signals of N_2O_5 and NO_3 inferred at 62 amu in a TD-CIMS: chemical interference or a real atmospheric phenomenon?

X. Wang^{1,2}, T. Wang^{1,2}, C. Yan¹, Y. J. Tham¹, L. Xue^{1,2}, Z. Xu^{1,2}, and Q. Zha¹

¹Department of Civil and Environmental Engineering, The Hong Kong Polytechnic University, Hong Kong, China

²Environment Research Institute, Shandong University, Ji'nan, Shandong, China

Correspondence to: T. Wang (cetwang@polyu.edu.hk)

Received: 18 July 2013 – Published in Atmos. Meas. Tech. Discuss.: 19 August 2013

Revised: 23 November 2013 – Accepted: 25 November 2013 – Published: 6 January 2014

Abstract. Dinitrogen pentoxide (N_2O_5) and the nitrate radical (NO_3) play important roles in atmospheric chemistry, yet accurate measurements of their concentrations remain challenging. A thermal dissociation chemical ionization mass spectrometer (TD-CIMS) was deployed to an urban site in Hong Kong to measure the sum of N_2O_5 and NO_3 in autumn 2010 based on the signals of NO_3^- at 62 amu which has also been adopted in previous studies reported in literature. To our surprise, very large signals of $\text{N}_2\text{O}_5 + \text{NO}_3$ were frequently observed at 62 amu in the daytime, with equivalent $\text{N}_2\text{O}_5 + \text{NO}_3$ mixing ratios in the range of 200–1000 pptv. To investigate this unusual phenomenon, various interference tests and measurements with different instrument configuration were conducted. It was found that peroxy acetyl nitrate (PAN) contributed to measurable signals at 62 amu, and more importantly, this interference increased significantly with co-existence of NO_2 . Nitric acid (HNO_3), on the other hand, had little interference to the detection of $\text{N}_2\text{O}_5/\text{NO}_3$ via the NO_3^- ion in our TD-CIMS. According to the test results, the interference from PAN and NO_2 could have contributed to 30–50 % of the average daytime (12:00–16:00, local time) $\text{N}_2\text{O}_5 + \text{NO}_3$ signal at our site. On the other hand, evidence exists for the presence of elevated daytime N_2O_5 , in addition to the daytime signal at 62 amu. This includes (1) daytime N_2O_5 measured via the $\text{I}(\text{N}_2\text{O}_5)^-$ cluster ion with an unheated inlet, which was subjected to minimum interferences, and (2) observation of elevated daytime ClNO_2 (a product of N_2O_5 hydrolysis) during a follow-up study. In view of the difficulty in accurately quantifying the contribution from the interferences of PAN and NO_2 and untested potential

interfering chemicals in the real atmosphere, we caution the use of 62 amu in the TD-CIMS for measuring ambient N_2O_5 in a high NO_x environment like Hong Kong. Additional studies are needed to re-examine the daytime issue using other measurement techniques.

1 Introduction

The nitrate radical (NO_3) and dinitrogen pentoxide (N_2O_5) play important roles in the nocturnal tropospheric chemistry. NO_3 is among the most important oxidants in the atmosphere, particularly for biogenic hydrocarbons and sulfur-containing compounds (Atkinson, 1991). N_2O_5 has long been recognized as a key intermediate in the transformation of nitrogen oxides ($\text{NO}_x = \text{NO} + \text{NO}_2$) to aerosol nitrates (Riemer et al., 2003; Aldener et al., 2006; Chang et al., 2011; Brown and Stutz, 2012). Recent studies have also demonstrated an important role of N_2O_5 hydrolysis at night in chlorine activation and the subsequent effect on the next-day's ozone formation (Osthoff et al., 2008; Simon et al., 2009; Thornton et al., 2010).

Due to the low ambient abundances and high reactivity, accurate measurements of atmospheric N_2O_5 and NO_3 have been challenging. Based on the strong absorption of NO_3 in the visible spectrum at 662 nm, several optical techniques have been developed to measure the ambient NO_3 , including long-path differential optical absorption spectroscopy (DOAS) (Platt et al., 1980; Atkinson et al., 1986), cavity ring-down spectroscopy (CRDS) (Brown et al., 2001, 2002),

laser-induced fluorescence (LIF) (Wood et al., 2003, 2005; Matsumoto et al., 2005), and cavity enhanced absorption spectroscopy (CEAS) (Venables et al., 2006; Langridge et al., 2008). N_2O_5 is determined by using a heated channel to decompose it into NO_3 or from the calculation according to the fast equilibrium between N_2O_5 with NO_3 and NO_2 .

Another technique for detecting ambient N_2O_5 and NO_3 is chemical ionization mass spectrometry (CIMS) which combines the ion-molecule chemistry with mass spectrometry detection. This technique was originally used in the laboratory to study the heterogeneous uptake kinetics of N_2O_5 and ClONO_2 (e.g., Hu and Abbatt, 1997; Thornton et al., 2003; Thornton and Abbatt, 2005), and later on was applied in field measurements (Slusher et al., 2004; Zheng et al., 2008). The fundamental of this method is the reaction of I^- (the reagent ion) with N_2O_5 (and/or NO_3) forming the NO_3^- ion that can be detected at 62 amu. Previous laboratory studies suggested that the $\text{N}_2\text{O}_5/\text{NO}_3$ measurement at 62 amu may be subject to interferences from other N-containing trace gases such as HNO_3 , HO_2NO_2 , ClONO_2 and others (see references listed in Table 1). On the other hand, a field inter-comparison of a TD-CIMS and a cavity ring-down system showed a high degree of correlation between the two methods (Huey, 2007; Chang et al., 2011) indicating the capability of this method in measuring N_2O_5 in the atmosphere. Although non-negligible and varying background signals at 62 amu have been observed during field studies (Chang et al., 2011), there have been no reports of detailed assessment of these potential interferences under different atmospheric conditions and for different configurations of CIMS. Kercher et al. (2009) developed a method to detect N_2O_5 via the $\text{I}(\text{N}_2\text{O}_5)^-$ cluster ion at 235 amu with an unheated inlet to address the interference at 62 amu in their CIMS, however this method has its own limitations including a lower sensitivity and larger impact of water vapor than that at 62 amu.

In autumn 2010, a TD-CIMS (thermal dissociation-CIMS), which is the same type used by Slusher et al. (2004) and Huey (2007), was deployed to an urban site in Hong Kong which is characterized by large quantities of NO_x , ozone and particulate matters. Unexpectedly, large signals of $\text{N}_2\text{O}_5 + \text{NO}_3$ inferred at 62 amu were frequently observed during the daytime in our study. To investigate this unusual observation, we have conducted a series of laboratory and field tests, including testing interferences individually and in combination from PAN, NO_2 , O_3 , HNO_3 , detection of N_2O_5 by using a cold inlet via detection of the $\text{I}(\text{N}_2\text{O}_5)^-$ cluster ion at 235 amu, and examination of daytime ClONO_2 which is a product of N_2O_5 hydrolysis. In this paper, we will first present the ambient observations, and then the detailed test results on interferences and other evidence for daytime N_2O_5 . A surprising result from the tests is that PAN+ NO_2 can contribute a large interference to the TD-CIMS signal at 62 amu, which has not been reported in previous studies. Despite large chemical interference, we show that the observed



Fig. 1. Location of the measurement site and surrounding layout.

daytime N_2O_5 signal may be in part due to a real contribution from NO_3 and N_2O_5 .

2 Experiment and methodology

2.1 Measurement site

The field study was conducted in the urban center of Hong Kong ($22^\circ 18' \text{N}$, $114^\circ 11' \text{E}$; $\sim 15 \text{ m}$ above sea level). The sampling site was located on the rooftop of a seven-story teaching building (about 20 m above the ground) on the campus of the Hong Kong Polytechnic University (PolyU). To the southeast of the campus, there is a cross-harbor tunnel in the north–south direction with a large flow of vehicles, especially during rush hours (see Fig. 1). Victoria Harbor is located about 1 km south of the measurement site. Thus emissions from vehicles and marine vessels are the most important local anthropogenic sources. The sampling site is surrounded by business districts, tourism and residential areas, with no large industrial sources nearby. The field measurements were carried out from 15 October to 4 December 2010, which is the season with the most severe photochemical pollution in Hong Kong (Wang et al., 2009).

2.2 The CIMS apparatus

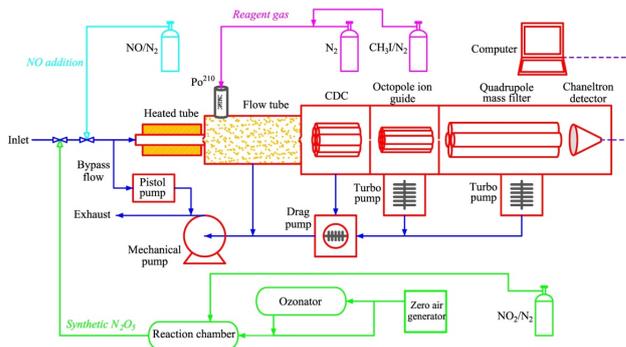
A TD-CIMS was deployed to measure the sum of N_2O_5 and NO_3 in this study. The system was developed at the *Georgia Institute of Technology* and is based on a soft and selective ionization process resulting from the reaction between a reagent ion and the target compounds, with the generated ions detected by a mass spectrometer. In the present study, the measurement method and operating parameters of TD-CIMS were the same as those described by Slusher et al. (2004) which was configured to simultaneously measure ambient N_2O_5 and PANs. The schematic diagram of our TD-CIMS is shown in Fig. 2. I^- , which was produced from passing a flow of 2 sccm of 0.3 % $\text{CH}_3\text{I}/\text{N}_2$ through an alpha ion source (Po-210), served as the reagent ion. Ambient air samples were drawn through a PFA-Teflon tube (I.D., 9.5 mm; O.D., 12.7 mm; length, 2 m) to the CIMS at a flow rate of 1.5 standard liters per minute (SLPM). To reduce the residence

Table 1. Ion-molecule reactions with product of NO_3^- and the rate constants.

| Reaction | Rate constant ($\text{molecule}^{-1} \text{cm}^3 \text{s}^{-1}$) | Comments | Reference |
|--|---|--------------------|----------------------------|
| $\text{I}^- + \text{HNO}_3 \rightarrow \text{NO}_3^- + \text{HI}$ | $< 5 \times 10^{-11}$ | | Fehsenfeld et al. (1975) |
| $\text{I}^- + \text{ClONO}_2 \rightarrow \text{NO}_3^- + \text{ICl}$ | 9×10^{-10} | $\pm 40\%$ | Huey et al. (1995) |
| $\text{I}^- + \text{BrONO}_2 \rightarrow \text{NO}_3^- + \text{IBr}$ | Unknown | | Hanson et al. (1996) |
| $\text{I}^- + \text{HO}_2\text{NO}_2 \rightarrow \text{NO}_3^- + \text{HIO}$ | Unknown | | Zhang et al. (1997) |
| $\text{CO}_3^- + \text{HNO}_3 \rightarrow \text{NO}_3^- + \text{products}$ | 8×10^{-10} | | Fehsenfeld et al. (1975) |
| $\text{CO}_3^- + \text{NO}_2 \rightarrow \text{NO}_3^- + \text{CO}_2$ | 2×10^{-10} | $\pm 50\%$ | Ikezoe and Viggiano (1987) |
| $\text{CO}_4^- + \text{NO} \rightarrow \text{NO}_3^- + \text{CO}_2$ | 4.8×10^{-11} | $\pm 30\%$ | Ikezoe and Viggiano (1987) |
| $\text{C}_2\text{H}_2\text{N}^- + \text{HNO}_3 \rightarrow \text{NO}_3^- + \text{CH}_3\text{CN}$ | 1.4×10^{-9} | 297 K | Ikezoe and Viggiano (1987) |
| $\text{C}_2\text{H}_3\text{O}_2^- + \text{HNO}_3 \rightarrow \text{NO}_3^- + \text{C}_2\text{H}_4\text{O}_2$ | Unknown | | Roberts et al. (2010) |
| $\text{Cl}^- + \text{HNO}_3 \rightarrow \text{NO}_3^- + \text{HCl}$ | 1.6×10^{-9} | $\pm 40\%$ | Ikezoe and Viggiano (1987) |
| $\text{NO}_2^- + \text{HNO}_3 \rightarrow \text{NO}_3^- + \text{HONO}$ | 1.6×10^{-9} | $\pm 40\%$ | Ikezoe and Viggiano (1987) |
| $\text{NO}_2^- + \text{NO}_2 \rightarrow \text{NO}_3^- + \text{NO}$ | $< 2 \times 10^{-13}$ | | Ikezoe and Viggiano (1987) |
| $\text{NO}_2^- + \text{N}_2\text{O} \rightarrow \text{NO}_3^- + \text{N}_2$ | $< 1 \times 10^{-12}$ | | Ikezoe and Viggiano (1987) |
| $\text{NO}_2^- + \text{O}_3 \rightarrow \text{NO}_3^- + \text{O}_2$ | 1.2×10^{-10} | $\pm 40\%$ | Ikezoe and Viggiano (1987) |
| $\text{O}_3^- + \text{NO}_2 \rightarrow \text{NO}_3^- + \text{O}_2$ | 2.8×10^{-10} | $\pm 30\%$, 280 K | Ikezoe and Viggiano (1987) |

time of air samples in the sampling tube an extra bypass flow of 7.7 SLPM was directed to the exhaust. Immediately before the flow tube, the last 14.8 cm of inlet was heated to 180 °C. The temperature of the air on the axis of the heated inlet was estimated at ~ 117 °C (Slusher et al., 2004), and under these conditions more than 99 % of the N_2O_5 would decompose into NO_3 with an ambient NO_2 level of 40 ppbv. The produced and the original NO_3 then reacted with I^- to produce NO_3^- ions in the flow tube, which were subsequently detected and quantified by a quadruple mass spectrometer (Extrel 150-QC, with a FWHM resolution of 2000) at 62 amu with a temporal resolution of 6 s. The NO_3^- signal is proportional to the concentration of total NO_3 , the I^- signal, the effective reaction rate constant between them, and the effective reaction time (Huey, 2007). Figure 3 depicts mass spectra of ambient air in urban Hong Kong obtained both during the daytime and at night, which clearly shows the I^- signal at 127 amu, NO_3^- signal at 62 amu, $\text{CH}_3\text{C}(\text{O})\text{O}^-$ (PAN) signal at 59 amu, etc. The mass peak at 62 amu was well separated from that at 59 amu.

In the present study, the TD-CIMS instrument was calibrated once a week using the on-line N_2O_5 synthesis method (Bertram et al., 2009). The calibration source was generated from the reactions of NO_2 with O_3 and subsequently NO_3 with NO_2 . In a glass reaction chamber in a commercially available calibrator (Model 6100, EnviroNics) 90 sccm of 2.5 ppmv NO_2 in nitrogen gas was mixed with 480 sccm of 2.1–4.6 ppmv O_3 which was generated by UV photolysis of O_2 in zero air (Model 111, Thermo Environmental Instruments (TEI)). The reactions happened during a time period of one minute, and then the output was diluted to 6 slpm by zero air. The concentrations of the prepared N_2O_5

**Fig. 2.** Schematic diagram of the TD-CIMS deployed in this study.

were determined via the change in NO_2 concentrations after adding ozone, and conversely verified by the change in ozone after adding NO_2 . Zero air that was free from moisture and aerosols served as the diluent so as to prevent the hydrolysis of N_2O_5 during the calibrations (see Fig. 2). NO_2 was monitored with a chemiluminescence analyzer equipped with a photolytic converter which ensured measurements of the true NO_2 (Ryerson et al., 2000; Xu et al., 2013). When inputting 36.7 ppbv of NO_2 and 170.4 ppbv of O_3 , the generated N_2O_5 was 1.08 ppbv. Note that there could be slight overestimation in the concentration of N_2O_5 standard determined in our study, because a small fraction ($< 10\%$) of N_2O_5 could be converted into gaseous nitric acid (Bertram et al., 2009). The loss of total nitrogen in the calibration system was examined with a chemiluminescence NO_x analyzer equipped with a molybdenum converter by comparing the signal with

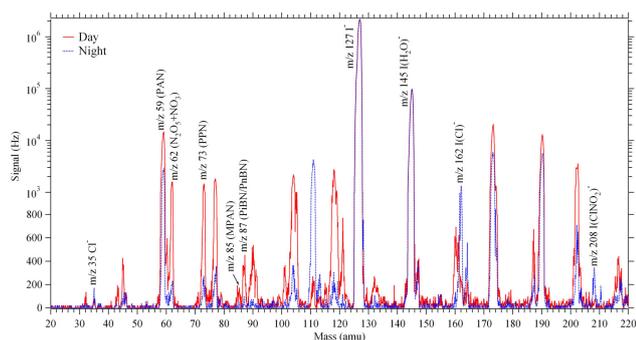


Fig. 3. Mass spectra of ambient air in urban Hong Kong at day and night by iodide TD-CIMS with heating the inlet tube to 180 °C.

and without adding O₃ to NO₂, which indicated insignificant loss.

Based on the relative standard deviation of the sample signal, the precision of our TD-CIMS was 3% for 1000 pptv N₂O₅. The sensitivity of N₂O₅ during the campaign was 2.8 ± 0.2 (mean \pm SD) Hz pptv⁻¹. The instrument background was automatically measured for 2 min once an hour by adding a small flow (5 mL min⁻¹) of NO (1000 ppm) to the sample flow (diluted to 9.2 SLPM), titrating NO₃ and thus N₂O₅. The background signal of the NO₃⁻ ion during this field study was 71.7 ± 36.0 (mean \pm SD) Hz. Figure 4 shows the signals of NO₃⁻ (at 62 amu) with hourly automatic background determinations on 30 November, when the equivalent maximum concentration of N₂O₅ + NO₃ was inferred. Clearly, the signals exhibited relatively low background compared to ambient N₂O₅ + NO₃ during both day and night. According to three times the standard deviation of the background signal, the typical detection limit of N₂O₅ for 6 s average time was estimated to be 39 pptv for our TD-CIMS.

2.3 Other instruments

In addition to the TD-CIMS, a large number of other instruments were deployed concurrently. Here, we briefly describe those that were used to aid the presentation of the N₂O₅ and NO₃ data. O₃ was measured by a commercial UV photometric analyzer (Model 49i, TEI, USA). NO and NO₂ were analyzed with a chemiluminescence instrument (Model 42i, TEI) equipped with a photolytic NO₂ converter (Air Quality Design, USA) (Xu et al., 2013). Solar radiation was measured using a LI-200 Pyranometer Sensor (LI-COR, USA). The ambient RH and temperature were monitored with a RH/temperature probe (Model 41382VC/VF, M.R. YOUNG, USA). During the field measurements, the minute-average data of trace gases and meteorological parameters were collected in real time by a data logger (Model 8816, Environmental Systems Corporation, USA).

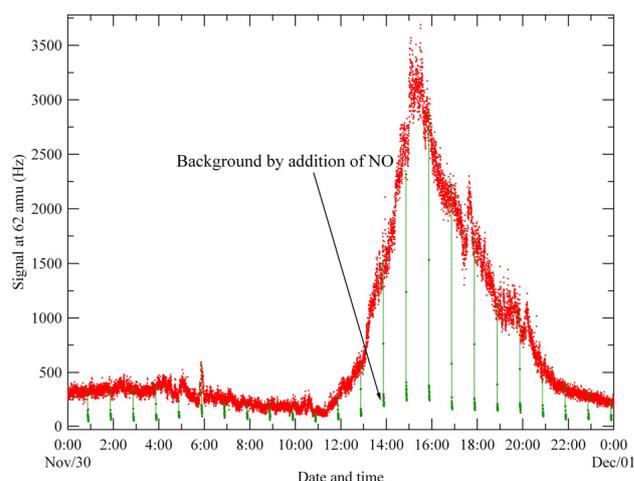


Fig. 4. Raw NO₃⁻ (62 amu) signal with hourly automatic background detection by adding excess NO, 30 November 2010.

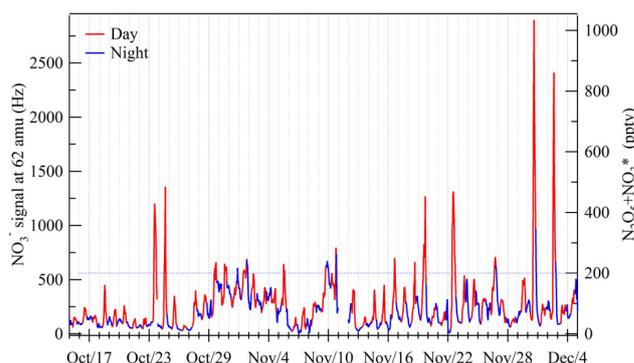


Fig. 5. Time series of hourly average NO₃⁻ signal at 62 amu and apparent N₂O₅ + NO₃ (N₂O₅ + NO₃^{*}) concentration measured in Hong Kong.

3 Observation results

The time series of hourly NO₃⁻ signals at 62 amu (after subtracting the background by adding NO) and apparent N₂O₅ + NO₃ mixing ratios (hereafter referred to as N₂O₅ + NO₃^{*}) from 15 October to 4 December 2010 is shown in Fig. 5. Similar to the measurement results obtained in other locations, several nighttime signal peaks of NO₃⁻ were noticed in our study (e.g., 31 October, 1 and 10 November). However, very large signals of N₂O₅ + NO₃ inferred by NO₃⁻ ion at 62 amu were frequently observed during the daytime (e.g., 23 and 24 October, 19, 22 and 30 November and 2 December). During the 50 day measurement period (excluding 11 November due to a lack of data), the average hourly apparent N₂O₅ + NO₃ concentration was 86.9 (\pm 85.6) pptv with the maximum value of 1033 pptv recorded at 15:00 LT (local time) on 30 November. The mean daytime apparent N₂O₅ + NO₃ mixing ratio was 102.5 pptv (06:00–17:59 LT),

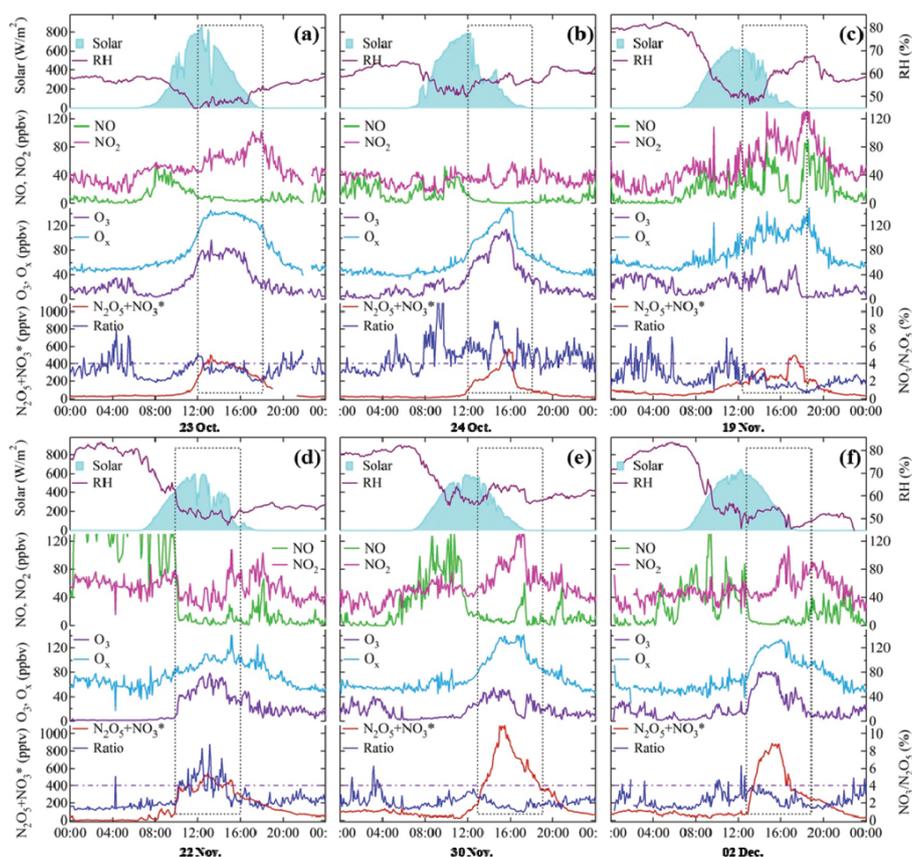


Fig. 6. Time series of apparent $\text{N}_2\text{O}_5 + \text{NO}_3$ ($\text{N}_2\text{O}_5 + \text{NO}_3^*$), $\text{NO}_3 / \text{N}_2\text{O}_5$ ratio, trace gases and meteorological parameters for six pollution episodes on (a) 23 October, (b) 24 October, (c) 19 November, (d) 22 November, (e) 30 November, and (f) 2 December 2010. The pollution episodes with daytime concentration peaks of N_2O_5 and NO_3 are shown in the panes.

even higher than the value of 71.3 pptv at nighttime (18:00–05:59 LT).

To learn more about the atmospheric conditions associated with the elevated daytime N_2O_5 and NO_3 signals at 62 amu, we present six cases during which the hourly apparent $\text{N}_2\text{O}_5 + \text{NO}_3$ values exceeded 400 pptv. Figure 6 presents the 5 min data of $\text{N}_2\text{O}_5 + \text{NO}_3^*$, O_3 , O_x ($\text{O}_3 + \text{NO}_2$), NO , NO_2 , RH , and solar radiation for these episodes. The ratio of $\text{NO}_3 / \text{N}_2\text{O}_5$, calculated based on the temperature dependent equilibrium among N_2O_5 , NO_3 and NO_2 , is also given. Inspection of the figure reveals that the daytime apparent $\text{N}_2\text{O}_5 + \text{NO}_3$ signal peaks appeared when both ozone and NO_2 were in high levels together with low levels of NO . However, daytime $\text{N}_2\text{O}_5 + \text{NO}_3$ concentrations calculated using both steady-state (Osthoff et al., 2006) and non-steady-state approaches (McLaren et al., 2010) were much lower (by a factor of 1–100) than observations. The daytime concentrations of apparent $\text{N}_2\text{O}_5 + \text{NO}_3$ in our study are also much higher than those of previous studies which reported daytime N_2O_5 and NO_3 concentrations of only a few to over 10 pptv (Geyer et al., 2003; Brown et al., 2005; Osthoff et al., 2006).

4 Chemical interferences

A major drawback of the TD-CIMS technique is that there could be potential interferences to the selected NO_3^- ion that can give rise to significant background noise at 62 amu (Chang et al., 2011). Previous laboratory studies have proposed many possible ion-molecule reactions yielding the NO_3^- ion, as summarized in Table 1. As stated earlier, the detailed results of these interferences in the real atmosphere have not been reported to date. To examine the influences of these potential interferences to our $\text{N}_2\text{O}_5 + \text{NO}_3$ measurements, a series of tests were conducted in various atmospheric conditions and in the laboratory during and after the measurement campaign.

The background determination by adding excess NO provided a first examination of the interferences. Such determined background reflects the interferences from compounds that are unreactive to NO , including most potential ones other than PAN and HO_2NO_2 . As stated above, the background was periodically measured once an hour throughout the campaign. A considerable background signal in our CIMS measurements at 62 amu was indeed seen

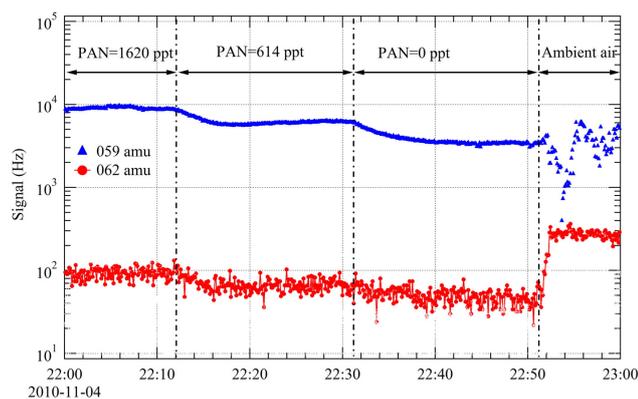


Fig. 7. Changes of signals at 59 and 62 amu when diluted synthetic PAN was introduced to the TD-CIMS.

during the present study, and they also exhibited a diurnal pattern with higher values in the late afternoon (see Fig. 4). This indicates that the $\text{N}_2\text{O}_5 + \text{NO}_3$ measurements via the 62 amu channel in our TD-CIMS were subject to some interference during the present study. However, the background signals were much lower than the total signal in ambient air. For the six cases with daytime $\text{N}_2\text{O}_5 + \text{NO}_3$ concentrations exceeding 400 pptv, the instrument background on average only accounted for 10 % of the ambient values. This background signal is due to most of the gases shown in Table 1 and has been accounted for (i.e., subtracted from the total signals at 62 amu) in our final data. The interference from PAN and HO_2NO_2 could not be determined because they also reacted with NO during zeroing. Their interferences are addressed below.

4.1 Interference of PAN

PAN is the most possible compound to interfere the TD-CIMS measurements in this study, not only due to its relatively high ambient abundances but also because it can escape from background determination by adding NO. In the previous studies of using a different technique of CRDS, higher N_2O_5 signals than the steady-state predictions were usually observed with elevated mixing ratios of PAN (Brown et al., 2005; Osthoff et al., 2006). The interference from PAN to the NO_3^- signals in our TD-CIMS was evaluated post field campaign by adding synthetic PAN to zero air and in the ambient air samples. The PAN was generated from a PAN calibrator (Meteorologie Consult GmbH), which is based on the reaction sequence of NO and acetone in ultrapure air with a Penray lamp. The concentrations of PAN that were added to the CIMS were quantified simultaneously by a NO_y analyzer (TEI 42CY). We also generated PAN using a conventional wet chemistry method by reacting peracetic acid with HNO_3 (Gaffney et al., 1984). The test results from the two PAN sources are consistent, thus this study only shows the result which used the PAN generated from the photolytic source.

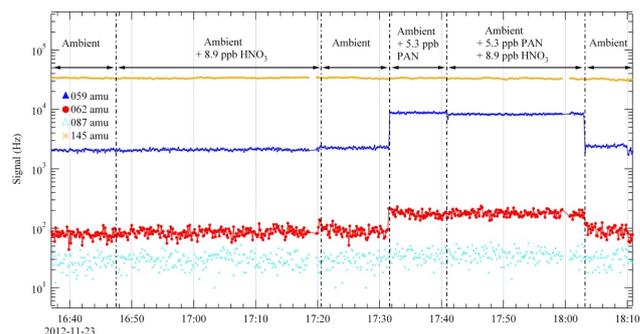


Fig. 8. Changes of raw signals of $\text{CH}_3\text{C}(\text{O})\text{O}^-$, NO_3^- , $\text{C}_3\text{H}_7\text{C}(\text{O})\text{O}^-$, $\text{I}(\text{H}_2\text{O})^-$ detected by I^- TD-CIMS when adding trace amounts of PAN, HNO_3 , or their mix to the ambient air.

For tests of PAN in zero air and in relatively clean ambient air at a coastal site (Hok Tsui), a measurable interference from PAN was observed at 62 amu. The result in zero air (i.e., the response of 62 amu signal to the PAN standard) indicates 23 ± 4 pptv of apparent $\text{NO}_3 + \text{N}_2\text{O}_5$ signal per ppbv of PAN (one example is shown in Fig. 7). Adding PAN to ambient air at the coastal site yielded similar results. Table 2 shows the results from four tests conducted with varying ambient pollution levels and meteorological conditions. Introducing 5.6–6.3 ppb of PAN to the ambient air resulted in an increase in the NO_3^- signal of 42–75 pptv of $\text{NO}_3 + \text{N}_2\text{O}_5$ in our TD-CIMS, again indicating interference from PAN to the field $\text{NO}_3/\text{N}_2\text{O}_5$ measurements via the NO_3^- by TD-CIMS. The mechanism for the interference of PAN on the NO_3^- signal is unclear. Some researchers proposed that the $\text{CH}_3\text{C}(\text{O})\text{O}^-$ ion (produced from the reaction of I^- with $\text{CH}_3\text{C}(\text{O})\text{O}_2$ – the product of PAN thermo-dissociation) likely reacts with HNO_3 to produce NO_3^- (Veres et al., 2008; Roberts et al., 2010). This appeared not to be the case in our CIMS as additions of both PAN and HNO_3 didn't lead to any increase at 62 amu compared to the addition of PAN alone (see Fig. 8).

A surprising finding was much larger interferences at 62 amu when the same spike tests were conducted at the PolyU site, which has very high NO_x concentrations, and the interference appeared to increase with ambient NO_2 suggesting that reaction between PAN and NO_2 leads to significant interference at 62 amu. To confirm this, a series of tests were conducted by mixing different levels of NO_2 with synthetic PAN (see an example in Fig. 9). Figure 10 shows the signal at 62 amu as a function of PAN and NO_2 concentrations in zero air, which clearly shows that the interference increases with both PAN and NO_2 . For example, at 5 ppbv of PAN, adding 60 ppbv of NO_2 produces a 400 ppt equivalent $\text{N}_2\text{O}_5/\text{NO}_3$ signal, compared to 150 pptv without NO_2 , indicating an amplifying effect of NO_2 on the previously reported PAN interference. The exact chemical reaction that leads to the interference is not clear. We believe that interference of PAN + NO_2 is related to the thermal dissociation of

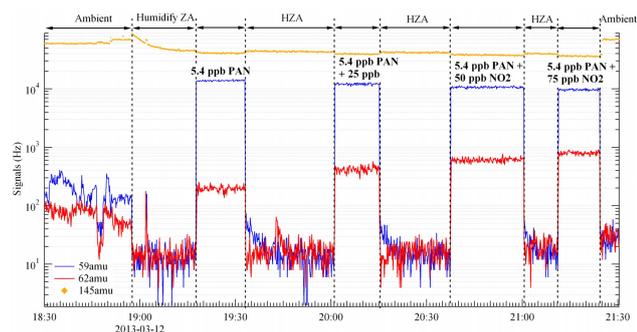


Fig. 9. Changes of signals at 59 and 62 amu when NO_2 was added to the sample flow of synthetic PAN.

PAN followed by radical reactions with NO_2 in the heated inlet. The radical reactions might serve as a source of NO_3 and N_2O_5 , or the products could react with I^- to produce NO_3^- ions which is detected at 62 amu. We will evaluate the possible contribution of PAN + NO_2 to the observed high daytime $\text{N}_2\text{O}_5 + \text{NO}_3$ for the 2010 autumn study in Sect. 5.

4.2 Interference of HNO_3

Another possible source of interference to the NO_3^- ion is HNO_3 since it has a NO_3 group. Several studies have proposed the production of NO_3^- from ion-molecule reactions involving HNO_3 , with somewhat inconsistent results. Huey and co-workers showed that the reaction of I^- with HNO_3 (producing NO_3^-) was rather inefficient. Thornton and coworkers on the other hand found a significant background signal (10–50 Hz) at the NO_3^- mass in a cold iodide CIMS under conditions of long ion-molecule reaction time, and attributed this to the reaction of HNO_3 with I^- (Thornton et al., 2003; Thornton and Abbatt, 2005). Roberts and coworkers reported that the HNO_3 was sensitive to the acetate ions and react to produce NO_3^- ions at the 62 amu (Veres et al., 2008; Roberts et al., 2010).

The relatively low background signals by adding NO in our study suggest insignificant interference from HNO_3 to our CIMS, which has been corrected during data reduction. This was further confirmed by the addition of HNO_3 . The test was carried out three times for varying ambient conditions to check the repeatability of the results, which are listed in Table 2 with an example being shown in Fig. 8. It can be seen that after introducing a trace amount of HNO_3 to the inlet tube of our TD-CIMS, there was no significant increase in the NO_3^- signal compared to the ambient air. Additionally, adding HNO_3 to a trace level of PAN which resulted in an acetate ion signal of $\sim 0.9 \times 10^4$ Hz, the NO_3^- signal showed no apparent change compared to that for only adding PAN. These results suggest that the HNO_3 itself and its mixture with acetate ions have no significant interference to the detection of $\text{N}_2\text{O}_5 + \text{NO}_3$ via the NO_3^- ion in our TD-CIMS. It should be noted that the result of non-reactivity

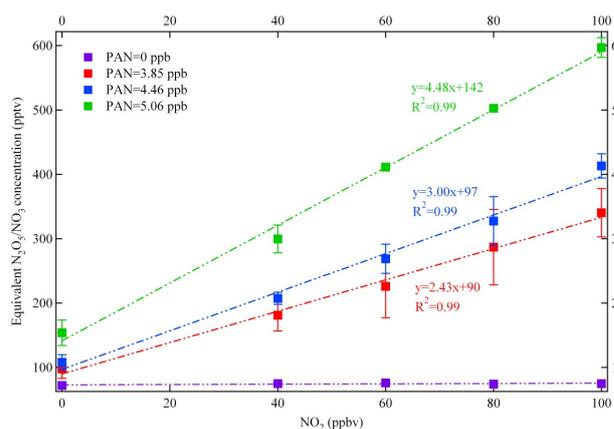


Fig. 10. Interference of NO_2 to equivalent NO_3^- concentration under different levels of PAN. Vertical bars are standard deviations of tests results.

of HNO_3 to the acetate ion in our TD-CIMS is different from that obtained by Roberts and coworkers using a NIPT-CIMS (negative-ion proton-transfer CIMS) (Veres et al., 2008; Roberts et al., 2010), possibly due to different configurations and operation conditions.

4.3 Interference of other possible compounds

ClONO_2 , BrONO_2 and HO_2NO_2 also react efficiently with I^- to produce NO_3^- (Huey et al., 1995; Hanson et al., 1996; Zhang et al., 1997; Amelynck et al., 2001). Again, the relatively low instrument background indicated small interference from ClONO_2 and BrONO_2 and any interference from them would have been corrected in our measurements because they cannot be removed by adding NO . The zeroing would not work for HO_2NO_2 . However, it is impossible for HO_2NO_2 to pass through the heated inlet tube in our TD-CIMS considering its thermally unstable nature.

In addition, the NO_3^- ion may also come from the ion-molecule reactions involving Cl^- (35 amu), $\text{C}_2\text{H}_2\text{N}^-$ (40 amu), NO_2^- (46 amu), O_3^- (48 amu), CO_3^- (60 amu), and CO_4^- (76 amu) (listed in Table 1). These reagent ions were observed in quite low levels (i.e., < 20 Hz for 40 and 48 amu, and 10–500 Hz for 35, 46, 60 and 76 amu) in our TD-CIMS during the field measurements (see Fig. 3). Therefore, the ion-molecule reactions induced by Cl^- , $\text{C}_2\text{H}_2\text{N}^-$, NO_2^- , O_3^- , CO_3^- and CO_4^- are also believed to have no significant influence to the NO_3^- detection based on the relationship between the product ion and reactants.

In summary, according to the above tests and discussions, reactions between PAN and NO_2 in the heated inlet are found to have significant interference to signal at 62 amu in our TD-CIMS, while HNO_3 and other compounds are not believed to have contributed to signal at 62 amu, although our tests are not exhaustive for including all other chemicals.

Table 2. Summary of the interference tests for the TD-CIMS.

| Test | Ambient conditions | | | | Standard addition* | | CIMS response | |
|------|--------------------------|-------------------------|----------|-----------|--------------------|----------------------------|---|---|
| | NO _y (ppb) | O ₃ (ppb) | T (°) | RH (%) | ΔPAN (ppb) | ΔHNO ₃ (ppb) | ΔN ₂ O ₅ (ppt) | ΔS ₆₂ /ΔS ₅₉ (cps/cps) |
| 1 | 3.5 | 69 | 25 | 83 | 6.3 | 0 | 75 | 0.014 |
| 2 | 4.8 | 71 | 25 | 87 | 5.8 | 0 | 42 | 0.009 |
| 3 | 4.5 | 60 | 23 | 83 | 5.9 | 0 | 69 | 0.014 |
| 4 | 14.0 | 35 | 22 | 69 | 5.6 | 0 | 62 | 0.013 |
| 5 | 7.2 | 70 | 24 | 81 | 0 | 1.4 | 1 | n.a. |
| 6 | 6.2 | 76 | 24 | 64 | 0 | 9.1 | 1 | n.a. |
| 7 | 15.2 | 38 | 22 | 69 | 0 | 8.9 | 2 | n.a. |
| 8 | 12.8 | 41 | 21 | 68 | 5.3 | 8.9 | 62 | n.a. |

* The concentrations of PAN and HNO₃ that were added to the ambient air were determined by a NO_y analyzer.

5 Contribution of interference to daytime concentrations

As indicated above, laboratory and field tests revealed significant interference from PAN and NO₂ to 62 amu in our TD-CIMS. The contribution from this interference to the observed daytime N₂O₅ + NO₃ in this study was examined and corrected based on the tests results in Fig. 10. During the field campaign at PolyU, PAN was concurrently measured by the same TD-CIMS and a GC-ECD instrument (gas chromatography with an electron capture detector). For the six episodes shown in Fig. 6, the interference would contribute to 41.3–67.0 %, 29.0–39.8 %, 32.2–73.2 %, 25.6–49.2 %, 19.4–43.4 %, 20.4–77.8 % of the daytime signals in the 6 episodes (two examples on 23 October and 30 November are shown in Fig. 11). The remaining daytime signal after correction may be a real contribution from N₂O₅ and NO₃, but interferences from other untested chemicals (e.g., organic nitrates in addition to PAN) cannot be ruled out. As interferences at the 62 amu channel are large and variable, and some of them may not have been identified in our tests, it is difficult to determine the fraction of real signal from N₂O₅ + NO₃ in our study at 62 amu.

To compare the observed signals with the concentrations predicted by known chemistry, Fig. 12 shows the source strength and loss frequency of NO₃, the predicted N₂O₅ + NO₃ levels for the six polluted episodes, together with corrected N₂O₅ + NO₃ data (i.e., subtracting the interference due to PAN + NO₂). In brief, the reaction of NO₃ with NO is the dominant loss pathway at the study site. In most cases, the NO concentration was relatively high, and the calculated values using steady-state assumption (see SI) were much lower than the observed values (after correction for interference due to PAN + NO₂). There is one exception in the late afternoon on 24 October. At 16:00 on this day when O₃, NO₂ and NO concentrations were 96.1, 49.1 and 0.3 ppbv, respectively, the predicted NO₃ and N₂O₅

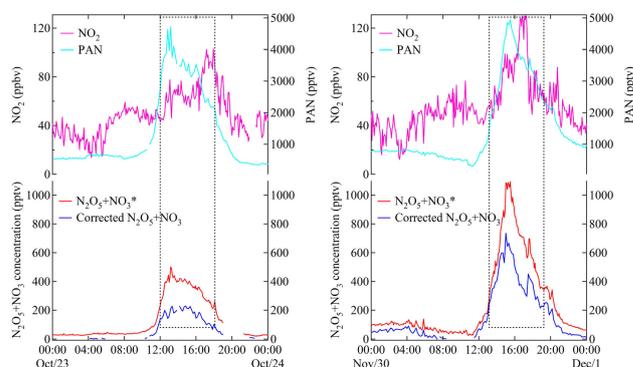


Fig. 11. Time series of corrected (N₂O₅+NO₃^{*}) and uncorrected N₂O₅ + NO₃, PAN, NO₂ and NO on 23 October and 30 November 2010.

concentration (5 min average) could reach 321 pptv if all major removal pathways are considered. In comparison, the measured N₂O₅ + NO₃ concentration was 286.5 pptv (with correction for PAN + NO₂ interference). The agreement between observation and prediction under the condition of very low NO in this case indicates that the elevated N₂O₅ in daytime is possible at the study site. The significant discrepancy in most cases may be attributed to other unknown measurement interference (e.g., chemical interference from organic nitrates in addition to PAN), unknown NO₃ source(s) (e.g., chemical reactions between Criegee intermediates and NO₂ suggested by recent studies (Welz et al., 2012; Ouyang et al., 2013), and invalidity of the steady-state assumption for air masses receiving fresh injection of NO). Further studies are required.

6 Other evidence of daytime NO₃ + N₂O₅ in Hong Kong

To check the measurement results at 62 amu, we attempted to measure ambient N₂O₅ with a cold version of CIMS

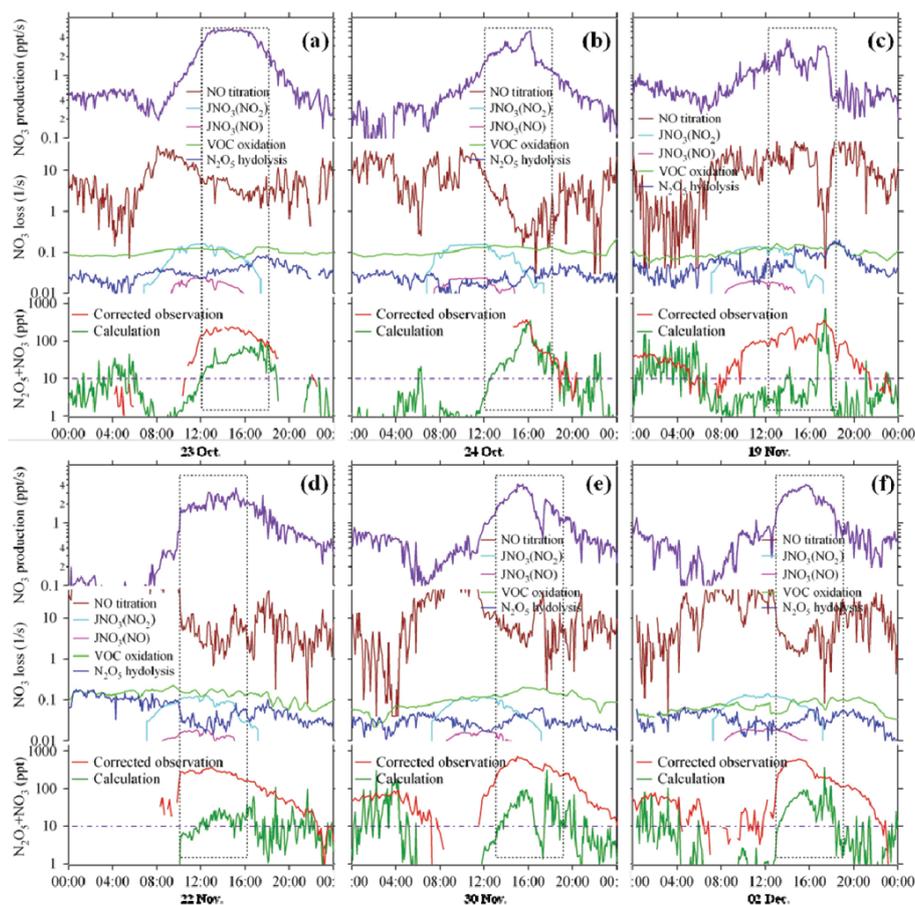


Fig. 12. NO_3 production rate from NO_2 and ozone, loss frequency of NO_3 from NO reaction, photolysis, VOC oxidation and N_2O_5 hydrolysis, steady-state calculated $\text{N}_2\text{O}_5 + \text{NO}_3$ and observed $\text{N}_2\text{O}_5 + \text{NO}_3$ after correction (for the interference from $\text{PAN} + \text{NO}_2$) for the six pollution episode days.

immediately after the present campaign at the same site. The CIMS was configured similar to that described by Kercher et al. (2009). When an unheated inlet tube is used, the reactions of N_2O_5 with I^- produce both the NO_3^- and $\text{I}(\text{N}_2\text{O}_5)^-$ cluster ions. The $\text{I}(\text{N}_2\text{O}_5)^-$ ion (at 235 amu) is thought to be free from the chemical interferences that can perturb the NO_3^- ion, and thus provides a better measure of N_2O_5 (Kercher et al., 2009). However, the $\text{I}(\text{N}_2\text{O}_5)^-$ sensitivity is influenced by the abundance of the water content in the sample flow. Similar to the setup with the heated inlet and the detection at 62 amu by TD-CIMS, signal at 235 amu with the cold CIMS also showed a daytime peak. Figure 13 gives an example of the ambient $\text{I}(\text{N}_2\text{O}_5)^-$ signal (in Hz) and related trace gases taken on 20 December 2010. The 235 amu signal exhibited generally good correlation with the 62 amu signal during both daytime and nighttime (figure not shown here). Calibration of the cold CIMS with the N_2O_5 standard in dry synthetic air gave a sensitivity of $0.55 \pm 0.003 \text{ Hz pptv}^{-1}$. But the true sensitivity for ambient air with moderately high humidity during study is believed to be larger, possibly by a factor of two or more, according to our recent lab tests.

Another independent piece of evidence for the daytime $\text{N}_2\text{O}_5 + \text{NO}_3$ in Hong Kong is concurrent increase in the mixing ratios of ClNO_2 (a product of N_2O_5 hydrolysis) observed in a follow-up study in western Hong Kong (Tung Chung; see SI for the experiment information). At this site, elevated $\text{N}_2\text{O}_5 + \text{NO}_3$ signals were also found at 62 amu at daytime during photochemical episodes, with the ClNO_2 signals (208 amu) showing concurrent increases. Figure 14 shows an example for 28 August 2011. On that day, apparent $\text{N}_2\text{O}_5 + \text{NO}_3$ inferred from 62 amu exhibited an afternoon peak of 670 pptv (uncorrected 5 min average), and ClNO_2 had a concurrent enhancement to 120 pptv. To estimate the levels of N_2O_5 that would be needed to sustain such amount of ClNO_2 , we assumed a photostationary steady state for ClNO_2 in the afternoon with an uptake coefficient of 0.03 for N_2O_5 hydrolysis on aerosol surfaces and a ClNO_2 yield of 10% with consideration of relatively low content of sea salt in aerosols at the site. The photolysis rate of ClNO_2 was estimated as $7.8 \times 10^{-4} \text{ s}^{-1}$ using the method by Simon et al. (2009), and the aerosol surface area was $979 \text{ mm}^2 \text{ m}^{-3}$ from concurrent measurements of aerosol size

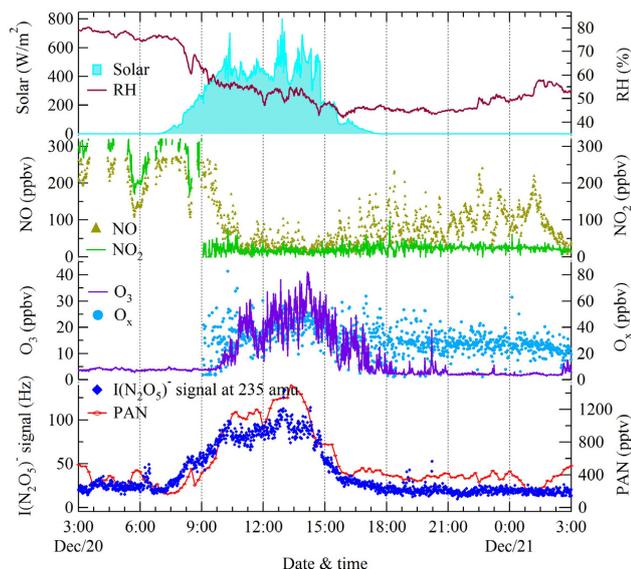


Fig. 13. Time series of $I(\text{N}_2\text{O}_5)^-$ signal at 235 amu, solar radiation and relative humidity in urban Hong Kong on 20 December 2010.

distribution. The calculation showed that to produce the observed 120 pptv of ClNO_2 , at least 518 pptv of N_2O_5 would be required. This result, although is not an absolute proof, supports the possibility of elevated daytime concentrations of $\text{N}_2\text{O}_5 + \text{NO}_3$ in Hong Kong. On this day the peak concentration of ClNO_2 was ~ 600 pptv at night and the ClNO_2 concentration exhibited good correlation ($R^2 = 0.78$) with the observed NO_3^- signals at 62 amu (figures not shown here).

7 Conclusions

A TD-CIMS technique, which has been previously applied to field measurements in the US, was deployed to measure ambient NO_3 and N_2O_5 in urban Hong Kong in a photochemical season. Surprisingly, large signals of $\text{NO}_3 + \text{N}_2\text{O}_5$ inferred from 62 amu were frequently observed in daytime, which is in contrast to our current understanding of reactive nitrogen chemistry. Subsequent laboratory and field tests provide new insights into chemical interferences in the TD-CIMS. In particular, we have discovered that reaction between NO_2 and PAN can amplify the interference from PAN at 62 amu. This interference could have contributed 30–50% to the average daytime NO_3 and N_2O_5 at our site. It is possible that the additional interferences from untested chemicals may have also contributed to the daytime signal at 62 amu. On the other hand, evidence exists to suggest that the elevated daytime signals at 62 amu may be in part due to real contribution from NO_3 or N_2O_5 . In viewing of the large and variable interferences at 62 amu and difficulty in correcting them, we conclude that it is not suitable to use the TD-CIMS to measure NO_3 or N_2O_5 at 62 amu in a high NO_x environment like the present study site. Adoption of either 235 amu with a cold

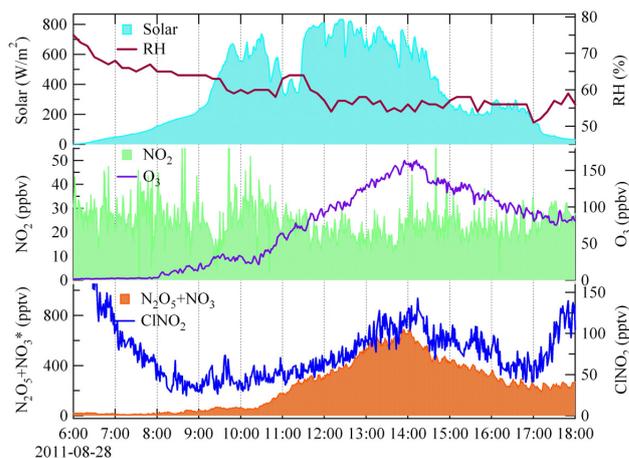


Fig. 14. A case showing concurrent increase of ClNO_2 and N_2O_5 at daytime during a photochemical episode at a suburban site (Tung Chung) in western Hong Kong.

inlet or an optics-based technique is recommended. We also suggest more studies to examine the abundance of daytime NO_3 and N_2O_5 in similar urban areas with co-existence of high ozone and NO_2 .

Supplementary material related to this article is available online at <http://www.atmos-meas-tech.net/7/1/2014/amt-7-1-2014-supplement.pdf>.

Acknowledgements. We thank Steven Poon and K.S. Lam for their help in the field study and setting up the instruments. This research was supported by the Environment and Conservation Fund of Hong Kong (project no. 7/2009), National Natural Science Foundation of China (41275123), Research Grants Council of Hong Kong (PolyU 5015/12P), and Niche Area Development Scheme of the Hong Kong Polytechnic University (1-BB94).

Edited by: E. C. Apel

References

- Aldener, M., Brown, S., Stark, H., Williams, E., Lerner, B., Kuster, W., Goldan, P., Quinn, P., Bates, T., and Fehsenfeld, F.: Reactivity and loss mechanisms of NO_3 and N_2O_5 in a polluted marine environment: Results from in situ measurements during New England Air Quality Study 2002, *J. Geophys. Res.*, 111, D23S73, doi:10.1029/2006JD007252, 2006.
- Amelynck, C., Stépien, C., Schoon, N., Catoire, V., Labonnette, D., Arijs, E., and Poulet, G.: Gas phase reactions of negative ions with ClONO_2 , *Int. J. Mass. Spectrom.*, 207, 205–215, doi:10.1016/S1387-3806(01)00373-6, 2001.
- Atkinson, R.: Kinetics and Mechanisms of the Gas-Phase Reactions of the NO_3 Radical with Organic Compounds, *J. Phys. Chem. Ref. Data*, 20, 459–507, doi:10.1063/1.555887, 1991.

- Atkinson, R., Winer, A. M., and Pitts, J. N.: Estimation of nighttime N_2O_5 concentrations from ambient NO_2 and NO_3 radical concentrations and the role of N_2O_5 in night-time chemistry, *Atmospheric Environment* (1967), 20, 331–339, doi:10.1016/0004-6981(86)90035-1, 1986.
- Bertram, T. H., Thornton, J. A., and Riedel, T. P.: An experimental technique for the direct measurement of N_2O_5 reactivity on ambient particles, *Atmos. Meas. Tech.*, 2, 231–242, doi:10.5194/amt-2-231-2009, 2009.
- Brown, S. S. and Stutz, J.: Nighttime radical observations and chemistry, *Chem. Soc. Rev.*, 41, 6405–6447, doi:10.1039/c2cs35181a, 2012.
- Brown, S., Stark, H., Ciciora, S., and Ravishankara, A.: In situ measurement of atmospheric NO_3 and N_2O_5 via cavity ring down spectroscopy, *Geophys. Res. Lett.*, 28, 3227–3230, doi:10.1029/2001GL013303, 2001.
- Brown, S. S., Stark, H., Ciciora, S. J., McLaughlin, R. J., and Ravishankara, A. R.: Simultaneous in situ detection of atmospheric NO_3 and N_2O_5 via cavity ring-down spectroscopy, *Rev. Sci. Instrum.*, 73, 3291–3301, doi:10.1063/1.1499214, 2002.
- Brown, S. S., Osthoff, H. D., Stark, H., Dubé, W. P., Ryerson, T. B., Warneke, C., de Gouw, J. A., Wollny, A. G., Parrish, D. D., and Fehsenfeld, F. C.: Aircraft observations of daytime NO_3 and N_2O_5 and their implications for tropospheric chemistry, *J. Photoch. Photobiol. A*, 176, 270–278, doi:10.1016/j.jphotochem.2005.10.004, 2005.
- Chang, W. L., Bhave, P. V., Brown, S. S., Riemer, N., Stutz, J., and Dabdub, D.: Heterogeneous Atmospheric Chemistry, Ambient Measurements, and Model Calculations of N_2O_5 : A Review, *Aerosol Sci. Tech.*, 45, 665–695, doi:10.1080/02786826.2010.551672, 2011.
- Fehsenfeld, F., Howard, C., and Schmeltekopf, A.: Gas phase ion chemistry of HNO_3 , *J. Phys. Chem.*, 63, 2835–2841, doi:10.1063/1.431722, 1975.
- Gaffney, J. S., Fajer, R., and Senum, G. I.: An improved procedure for high purity gaseous peroxyacyl nitrate production: Use of heavy lipid solvents, *Atmos. Environ.*, 18, 215–218, doi:10.1016/0004-6981(84)90245-2, 1984.
- Geyer, A., Aliche, B., Ackermann, R., Martinez, M., Harder, H., Brune, W., di Carlo, P., Williams, E., Jobson, T., and Hall, S.: Direct observations of daytime NO_3 : Implications for urban boundary layer chemistry, *J. Geophys. Res.*, 108, 4368, doi:10.1029/2002JD002967, 2003.
- Hanson, D. R., Ravishankara, A. R., and Lovejoy, E. R.: Reaction of BrONO_2 with H_2O on submicron sulfuric acid aerosol and the implications for the lower stratosphere, *J. Geophys. Res.*, 101, 9063–9069, doi:10.1029/96jd00347, 1996.
- Hu, J. H. and Abbatt, J. P. D.: Reaction Probabilities for N_2O_5 Hydrolysis on Sulfuric Acid and Ammonium Sulfate Aerosols at Room Temperature, *J. Phys. Chem. A*, 101, 871–878, doi:10.1021/jp9627436, 1997.
- Huey, L.: Measurement of trace atmospheric species by chemical ionization mass spectrometry: Speciation of reactive nitrogen and future directions, *Mass. Spectrom. Rev.*, 26, 166–184, doi:10.1002/mas.20118, 2007.
- Huey, L., Hanson, D., and Howard, C.: Reactions of SF_6^- and I^- with atmospheric trace gases, *J. Phys. Chem.*, 99, 5001–5008, doi:10.1021/j100014a021, 1995.
- Ikezo, Y. and Viggiano, A.: Gas phase ion-molecule reaction rate constants through 1986, Ion Reaction Research Group of the Mass Spectroscopy Society of Japan: Distributed by Maruzen Co., Tokyo, Japan, 1987.
- Kercher, J. P., Riedel, T. P., and Thornton, J. A.: Chlorine activation by N_2O_5 : simultaneous, in situ detection of ClNO_2 and N_2O_5 by chemical ionization mass spectrometry, *Atmos. Meas. Tech.*, 2, 193–204, doi:10.5194/amt-2-193-2009, 2009.
- Langridge, J. M., Ball, S. M., Shillings, A. J. L., and Jones, R. L.: A broadband absorption spectrometer using light emitting diodes for ultrasensitive, in situ trace gas detection, *Rev. Sci. Instrum.*, 79, 123110, doi:10.1063/1.3046282, 2008.
- Matsumoto, J., Kosugi, N., Imai, H., and Kajii, Y.: Development of a measurement system for nitrate radical and dinitrogen pentoxide using a thermal conversion/laser-induced fluorescence technique, *Rev. Sci. Instrum.*, 76, 064101–064111, doi:10.1063/1.1927098, 2005.
- McLaren, R., Wojtal, P., Majonis, D., McCourt, J., Halla, J. D., and Brook, J.: NO_3 radical measurements in a polluted marine environment: links to ozone formation, *Atmos. Chem. Phys.*, 10, 4187–4206, doi:10.5194/acp-10-4187-2010, 2010.
- Osthoff, H. D., Sommariva, R., Baynard, T., Pettersson, A., Williams, E. J., Lerner, B. M., Roberts, J. M., Stark, H., Goldan, P. D., and Kuster, W. C.: Observation of daytime N_2O_5 in the marine boundary layer during New England Air Quality Study–Intercontinental Transport and Chemical Transformation 2004, *J. Geophys. Res.*, 111, D23S14, doi:10.1029/2006JD007593, 2006.
- Osthoff, H. D., Roberts, J. M., Ravishankara, A. R., Williams, E. J., Lerner, B. M., Sommariva, R., Bates, T. S., Coffman, D., Quinn, P. K., Dibb, J. E., Stark, H., Burkholder, J. B., Talukdar, R. K., Meagher, J., Fehsenfeld, F. C., and Brown, S. S.: High levels of nitryl chloride in the polluted subtropical marine boundary layer, *Nat. Geosci.*, 1, 324–328, doi:10.1038/ngeo177, 2008.
- Ouyang, B., McLeod, M. W., Jones, R. L., and Bloss, W. J.: NO_3 radical production from the reaction between the Criegee intermediate CH_2OO and NO_2 , *Phys. Chem. Chem. Phys.*, 15, 17070–17075, doi:10.1039/c3cp53024h, 2013.
- Platt, U., Perner, D., Winer, A. M., Harris, G. W., and Pitts Jr., J. N.: Detection of NO_3 in the polluted troposphere by differential optical absorption, *Geophys. Res. Lett.*, 7, 89–92, doi:10.1029/GL007i001p00089, 1980.
- Riemer, N., Vogel, H., Vogel, B., Schell, B., Ackermann, I., Kessler, C., and Hass, H.: Impact of the heterogeneous hydrolysis of N_2O_5 on chemistry and nitrate aerosol formation in the lower troposphere under photo-smog conditions, *J. Geophys. Res.*, 108, 4144, doi:10.1029/2002JD002436, 2003.
- Roberts, J. M., Veres, P., Warneke, C., Neuman, J. A., Washenfelder, R. A., Brown, S. S., Baasandorj, M., Burkholder, J. B., Burling, I. R., Johnson, T. J., Yokelson, R. J., and de Gouw, J.: Measurement of HONO, HNCO, and other inorganic acids by negative-ion proton-transfer chemical-ionization mass spectrometry (NI-PT-CIMS): application to biomass burning emissions, *Atmos. Meas. Tech.*, 3, 981–990, doi:10.5194/amt-3-981-2010, 2010.
- Ryerson, T., Williams, E., and Fehsenfeld, F.: An efficient photolysis system for fast-response NO_2 measurements, *J. Geophys. Res.*, 105, 26447–26461, doi:10.1029/2000JD900389, 2000.
- Simon, H., Kimura, Y., McGaughey, G., Allen, D. T., Brown, S. S., Osthoff, H. D., Roberts, J. M., Byun, D., and Lee, D.: Modeling

- the impact of ClNO₂ on ozone formation in the Houston area, *J. Geophys. Res.*, 114, D00F03, doi:10.1029/2008jd010732, 2009.
- Slusher, D. L., Huey, L. G., Tanner, D. J., Flocke, F. M., and Roberts, J. M.: A thermal dissociation-chemical ionization mass spectrometry (TD-CIMS) technique for the simultaneous measurement of peroxyacyl nitrates and dinitrogen pentoxide, *J. Geophys. Res.*, 109, D19315, doi:10.1029/2004jd004670, 2004.
- Thornton, J. and Abbatt, J.: N₂O₅ reaction on submicron sea salt aerosol: Kinetics, products, and the effect of surface active organics, *J. Phys. Chem. A*, 109, 10004–10012, doi:10.1021/jp054183t, 2005.
- Thornton, J., Braban, C., and Abbatt, J.: N₂O₅ hydrolysis on submicron organic aerosols: the effect of relative humidity, particle phase, and particle size, *Phys. Chem. Chem. Phys.*, 5, 4593–4603, doi:10.1039/B307498F, 2003.
- Thornton, J., Kercher, J., Riedel, T., Wagner, N., Cozic, J., Holloway, J., Dub , W., Wolfe, G., Quinn, P., and Middlebrook, A.: A large atomic chlorine source inferred from mid-continental reactive nitrogen chemistry, *Nature*, 464, 271–274, doi:10.1038/nature08905, 2010.
- Venables, D. S., Gherman, T., Orphal, J., Wenger, J. C., and Ruth, A. A.: High Sensitivity in Situ Monitoring of NO₃ in an Atmospheric Simulation Chamber Using Incoherent Broadband Cavity-Enhanced Absorption Spectroscopy, *Environ. Sci. Technol.*, 40, 6758–6763, doi:10.1021/es061076j, 2006.
- Veres, P., Roberts, J. M., Warneke, C., Welsh-Bon, D., Zahniser, M., Herndon, S., Fall, R., and de Gouw, J.: Development of negative-ion proton-transfer chemical-ionization mass spectrometry (NI-PT-CIMS) for the measurement of gas-phase organic acids in the atmosphere, *Int. J. Mass. Spectrom.*, 274, 48–55, doi:10.1016/j.ijms.2008.04.032, 2008.
- Wang, T., Wei, X. L., Ding, A. J., Poon, C. N., Lam, K. S., Li, Y. S., Chan, L. Y., and Anson, M.: Increasing surface ozone concentrations in the background atmosphere of Southern China, 1994–2007, *Atmos. Chem. Phys.*, 9, 6217–6227, doi:10.5194/acp-9-6217-2009, 2009.
- Welz, O., Savee, J. D., Osborn, D. L., Vasu, S. S., Percival, C. J., Shallcross, D. E., and Taatjes, C. A.: Direct Kinetic Measurements of Criegee Intermediate (CH₂OO) Formed by Reaction of CH₂I with O₂, *Science*, 335, 204–207, doi:10.1126/science.1213229, 2012.
- Wood, E., Wooldridge, P., Freese, J., Albrecht, T., and Cohen, R.: Prototype for in situ detection of atmospheric NO₃ and N₂O₅ via laser-induced fluorescence, *Environ. Sci. Technol.*, 37, 5732–5738, doi:10.1021/es034507w, 2003.
- Wood, E. C., Bertram, T. H., Wooldridge, P. J., and Cohen, R. C.: Measurements of N₂O₅, NO₂, and O₃ east of the San Francisco Bay, *Atmos. Chem. Phys.*, 5, 483–491, doi:10.5194/acp-5-483-2005, 2005.
- Xu, Z., Wang, T., Xue, L. K., Louie, P. K. K., Luk, C. W. Y., Gao, J., Wang, S. L., Chai, F. H., and Wang, W. X.: Evaluating the uncertainties of thermal catalytic conversion in measuring atmospheric nitrogen dioxide at four differently polluted sites in China, *Atmos. Environ.*, 76, 221–226, doi:10.1016/j.atmosenv.2012.09.043, 2013.
- Zhang, R., Leu, M. T., and Keyser, L. F.: Heterogeneous chemistry of HO₂NO₂ in liquid sulfuric acid, *J. Phys. Chem. A*, 101, 3324–3330, doi:10.1021/jp963321z, 1997.
- Zheng, J., Zhang, R., Fortner, E. C., Volkamer, R. M., Molina, L., Aiken, A. C., Jimenez, J. L., Gaeggeler, K., Dommen, J., Dusanter, S., Stevens, P. S., and Tie, X.: Measurements of HNO₃ and N₂O₅ using ion drift-chemical ionization mass spectrometry during the MILAGRO/MCMA-2006 campaign, *Atmos. Chem. Phys.*, 8, 6823–6838, doi:10.5194/acp-8-6823-2008, 2008.