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Exploring the amplified role of HCHO in the formation of HMS and O_3 during the co-occurring $PM_{2.5}$ and O_3 pollution in a coastal city of southeast China

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Abstract. To develop effective strategies for controlling both $PM_{2.5}$ and O_3 levels, it is crucial to understand their synergistic mechanisms and key precursors and the atmospheric physiochemical processes involved. In this study, a wintertime co-occurring O_3 and PM_{25} pollution event in a coastal city in southeast China was investigated based on high-time-resolution measurements of criteria air pollutants and chemical compositions of $PM_{2.5}$, and O_3 precursors, such as NO_x , HCHO, and volatile organic compounds (VOCs). The results of this study revealed the characteristics of positively correlated PM2.5 and MDA8 O3 concentrations, and an increase in atmospheric oxidation capacity (AOC) during the cold seasons. Strong correlations ($R^2 = 0.415 - 0.477$) between HCHO, Fe, Mn, and sulfate concentrations were observed, suggesting the influence of catalyzed oxidation processes in the coastal city. Through an observation-based model (OBM) analysis coupled with the Regional Atmospheric Chemistry Mechanism, version 2 (RACM2) and the Chemical Aqueous-Phase Radical Mechanism, version 3.0 (CAPRAM 3.0), we found that high concentrations of precursors (SO₂ and HCHO), high relative humidity, and moderately acidic pH conditions enhanced the heterogeneous formation of hydroxymethanesulfonate (HMS) in PM_{2.5}. Furthermore, by employing an OBM coupled to the Master Chemical Mechanism (OBM-MCM), we verified that disabling the HCHO mechanism could decrease daytime net O_3 production rates by reducing the production rates of $HO_2 + NO$. These results were consistent with the daily values of AOC, OH, HO₂, and RO₂ concentrations. This study contributes to a better understanding of the significance of HCHO in photochemical reactions and the formation of HMS in a coastal city.

1 Introduction

Air pollution, dominated by fine particulate matter (PM_{2.5}) and ground-level ozone (O₃), is an important global environmental issue linked to climate change and human health, including cardiovascular and respiratory illnesses and mortality (Xiao et al., 2022; Vohra et al., 2022). To decrease global air pollution and associated mortality, the World Health Organization recently updated its air quality guideline for annual PM_{2.5} exposure from 10 to 5 μ g m⁻³ and added the average O₃ concentrations of no more than 60 μ g m⁻³ during the peak season (WHO, 2021). To develop two-pollutant control strategies to decrease both PM_{2.5} and O₃, there is a need to understand the synergistic mechanisms and spatiotemporal delineation between them (Ivatt et al., 2022; Li et al., 2019b).

There are complex synergistic effects between PM_{2.5} and O₃, due to common precursors (e.g., NO_x and volatile organic compounds, VOCs), atmospheric physiochemical processes, and weather systems (Li et al., 2019a; Shao et al., 2022; Jia et al., 2023; Zhang et al., 2022; Qin et al., 2021; Qu et al., 2023). An increase in O₃ concentration will increase oxidizing substances, such as OH, H₂O₂, and RCHO, which promote the oxidation of SO₂, NO_x, and VOCs to secondary inorganic and organic components in PM_{2.5} (Feng et al., 2020; Lu et al., 2019). On the contrary, multiphase reactions occur on the surface of atmospheric particles, such as the uptake of HCHO and HO₂ absorption, thus affecting the formation of O₃ (Song et al., 2022; Lou et al., 2014).

Formaldehyde (HCHO) plays an important role in the photochemical reaction process and secondary aerosol formation (Kalashnikov et al., 2022; Ma et al., 2020; Song et al., 2021; Zong et al., 2021). Most studies have focused on pollution characteristics and sources, particle uptake of HCHO, and their impacts on atmospheric oxidation capacity (Liu et al., 2022b; Wu et al., 2023; Zhang et al., 2021). Recent studies have shown that HCHO can react with hydrogen peroxide (H_2O_2) to produce hydroxymethyl hydroperoxide (HMHP), which rapidly oxidizes dissolved sulfur dioxide (SO₂, aqueous phase (aq)) to sulfate (Dovrou et al., 2022). Meanwhile, HCHO reacts with dissolved SO₂ (aq) to produce hydroxymethanesulfonate (HMS, HOCH₂SO₃), which, upon oxidation with the hydroxyl radical (OH), forms sulfate (Ma et al., 2020; Moch et al., 2018, 2020). In total, atmospheric HCHO contributes to sulfate formation in PM_{2.5} by producing HO₂ radicals and HMHP or HMS (Wu et al., 2023; Dovrou et al., 2022; Campbell et al., 2022). However, these studies highlight the necessity for more observation research to obtain evidence of the contributions of HCHO to HMS formation. HMS is an important organosulfur compound in the atmosphere, not only in cloud and fog but also in atmospheric aerosols (Munger et al., 1986; Dixon and Aasen, 1999). The misidentification of HMS as inorganic sulfate caused the overestimation of the observed particulate sulfate (Ma et al., 2020; Dovrou et al., 2022). However, in subtropical coastal regions with apparent HCHO production, further studies are required to investigate the impacts of HCHO on the formation of HMS and RO_x radicals.

Xiamen, a coastal city in southeast China, has frequently experienced PM_{2.5} pollution (with low O₃ concentrations) in winter and O₃ pollution in spring and autumn (Hong et al., 2022; Wu et al., 2019). Our previous studies mainly focused on the pollution characteristics of $PM_{2.5}$ or O_3 in different seasons and their sources associated with anthropogenic emissions and the East Asian monsoon (Liu et al., 2020, 2022a; Hong et al., 2021). At the end of winter and the beginning of spring in 2022, an outbreak of co-occurring O₃ and PM_{2.5} pollution was observed in Xiamen. Therefore, it provided a unique opportunity to study the impacts of HCHO on the formation of HMS and RO_x radicals. In the coastal region, there is an apparent alternation of polluted and clean air masses from continental and ocean areas and a local geographical environment, including relatively high humidity, dense vegetation, and strong atmospheric oxidation capacity (Hu et al., 2022; Wu et al., 2020). The potential roles and mechanisms of HCHO in the photochemical reaction process and HMS formation would differ from those in megacities of China, such as the Beijing-Tianjin-Hebei (BTH) area, the Yangtze River Delta (YRD), and the Pearl River Delta (PRD). Based on the observation-based model (OBM) analysis, the objectives of this study are to (1) characterize the wintertime co-occurring O₃ and PM_{2.5} pollution process in a coastal city, (2) elaborate the influence of HCHO on the heterogeneous formation of HMS in $PM_{2.5}$, and (3) explore the mechanisms of HCHO on O₃ pollution and the photochemical reaction process.

2 Methods and materials

2.1 Study area

The monitoring site (Institute of Urban Environment, Chinese Academy of Sciences; 24.61° N, 118.06° E) is located in Xiamen, a coastal city in southeast China (Fig. S2 in the Supplement). It is situated in a subtropical monsoon climate, with an average temperature of $18.5 \,^{\circ}$ C and a relative humidity of $63.3 \,\%$ during the wintertime observation. In autumn and winter, cold and dry air masses move northward from inland, while in late spring and summer, the prevailing air masses are southerly, characterized by warm air temperatures and high humidity. The air-monitoring supersite is located on the rooftop of a building, surrounded by residential buildings, educational institutions, a commercial zone, and freeways. The downtown area of Xiamen, with a high population density and frequent traffic jams, is located south of the monitoring site.

2.2 Observation

Gas and aerosol species, O₃ precursors, the photolysis rate, and meteorological parameters were continuously measured online from 17 February to 17 March 2022. Hourly mass concentrations of PM2.5 and PM10 were measured using a tapered element oscillating microbalance (TEOM1405, Thermo Scientific Corp., MA, USA). NO/NO₂, SO₂, and O_3 were monitored using continuous gas analyzers (TEI 42*i*, 43i, and 49i, Thermo Scientific Corp., MA, USA). A HCHO analyzer (FMS-100, Focused Photonics Inc., Hangzhou, China) was used to measure gaseous HCHO based on the Hantzsch reaction, according to our previous method (Liu et al., 2022b). Water-soluble inorganic ions (WSIIs) in PM2.5 $(Cl^{-}, SO_{4}^{2-}, NO_{3}^{-}, Na^{+}, K^{+}, NH_{4}^{+}, Mg^{2+}, and Ca^{2+})$ were measured hourly using the Monitor for AeRosols and Gases in ambient Air (MARGA 2080; Metrohm Applikon B.V.; Delft, the Netherlands). Simultaneously, organic carbon (OC) and elemental carbon (EC) in PM25 were measured using an OC/EC analyzer (model RT-4; Sunset Laboratory Inc.; Tigard, USA). Black carbon (BC) was monitored using an aethalometer (AE31, Magee Scientific, USA) with a PM_{2.5} cutoff inlet. In addition, concentrations of 22 elements (Al, Si, S, Fe, K, Mn, Pb, Ca, Zn, Ba, V, Cu, Ni, As, Cr, Ag, Se, Br, Hg, Sn, Ti, and Sb) were measured using a multi-metal monitor (Xact[™] 625, Cooper Environmental Services, LLT; Portland, USA). Strict quality assurance and quality control procedures were applied, and the maintenance and accuracy of all online instruments were validated (Hong et al., 2021). The detection limits, time resolutions, and measured uncertainties of air pollutants are shown in Table S3.

A gas chromatograph with flame-ionization coupled with a mass spectrometer (GC-FID/MS, TH-300B, Wuhan, China) was used to measure ambient VOCs with 1h time resolution, following the method from our previous studies (Liu et al., 2020a, b). Briefly, the air sample was preconcentrated by cooling to -160 °C in a cryogenic trap, then heated to 100 °C, and subsequently transferred to the secondary trap using high-purity helium (He). The low-carbon (C_2-C_5) hydrocarbons were detected using a flame ionization detector (FID) with a PLOT (Al₂O₃/KCl) column $(15 \text{ m} \times 0.32 \text{ mm} \times 6.0 \mu\text{m})$, while other VOC species were quantified using a GC / MS with a DB-624 column (60 m \times $0.25 \text{ mm} \times 1.4 \mu\text{m}$). The instrument can quantify 106 VOC species, including 29 alkanes, 11 alkenes, 1 alkyne, 17 aromatics, 35 halogenated hydrocarbons, and 13 OVOCs (oxygenated VOCs). Calibration was performed daily at 23:00 LT (local time) using the standard mixtures of US EPA PAMS and TO-15. The detection limits of the measured VOCs ranged from 0.02 to 0.30 ppby.

Ambient meteorological parameters, including relative humidity (RH), temperature (T), wind speed (WS), and wind direction (WD), were obtained using an ultrasonic atmospherium (150WX, Airmar, USA). Photolysis frequencies and HCHO were measured using a photolysis spectrometer (PFS-100) and a formaldehyde monitor (FMS-100; Focused Photonics Inc., Hangzhou, China), respectively. The photolysis rate constants include $J(O^1D)$, $J(NO_2)$, $J(H_2O_2)$, J(HONO), J(HCHO), and $J(NO_3)$. The distribution of fire spots during the observation periods was obtained from the Fire Information for Resource Management System (https://firms.modaps.eosdis.nasa.gov/map/, last access: 21 July 2022). The data for boundary layer height (BLH) were obtained from the European Centre for Medium-Range Weather Forecasts (ECMWF) ERA5 hourly reanalysis dataset (https://www.ecmwf.int/en/forecasts/datasets/ reanalysis-datasets/era5, last access: 24 March 2023). The resolution of ERA-5 reanalysis is $0.25^{\circ} \times 0.25^{\circ}$.

2.3 Positive matrix factorization (PMF) analysis

The PMF 5.0 model was applied to quantify high-timeresolution sources of $PM_{2.5}$ during the observation periods. The details of the model analysis were described in our previous studies (Hong et al., 2021; Liu et al., 2020). Briefly, Eq. (1) demonstrates *j* compound species in the *i*th sample as the concentration from *p* independent sources.

$$x_{ij} = \sum_{k=1}^{p} g_{ik} f_{ki} + e_{ij},$$
(1)

where e_{ij} is the residual for each species, f_{kj} is the fraction of the *j*th species from the *k*th source, g_{ik} is the species contribution of the *k*th source to the *i*th sample, x_{ij} is the *j*th species concentration measured in the *i*th sample, and *p* is the total number of independent sources. The *Q* value (Eq. 2), based on the uncertainties (μ), was used to evaluate the steadiness of the solution.

$$Q = \sum_{i=1}^{n} \sum_{j=1}^{m} \left[\frac{x_{ij} - \sum_{k=1}^{p} g_{ik} f_{kj}}{\mu_{ij}} \right]^2$$
(2)

2.4 Observation-based model (OBM)

An OBM coupled to the Master Chemical Mechanism (OBM-MCM) is employed to simulate in situ atmospheric photochemical processes and quantify the O₃ production rate, AOC, and OH reactivity. Details of the OBM-MCM were reported in our previous studies (Liu et al., 2022a, b). In summary, monitoring data with a 1 h time resolution of air pollutants (i.e., O₃, CO, NO, NO₂, HONO, SO₂, and VOCs), meteorological parameters (i.e., T, P, and RH), and photolysis rate constants ($J(O^1D)$, $J(NO_2)$, $J(H_2O_2)$, J(HONO), J(HCHO), and $J(\text{NO}_3)$) were input into the OBM-MCM as constraints for the model simulation. The photolysis rates of other molecules were determined by solar zenith angle and scaled using measured JNO₂ values (Saunders et al., 2003). According to our previous studies, the model incorporates the physical process of deposition within the boundary layer height (BLH), which varies from 300 m during nighttime to 1500 m during the daytime in winter (Li et al., 2018; Liu et al., 2022a, b). Therefore, dry deposition velocities were used to simulate the deposition loss of certain reactants in the atmosphere (Zhang et al., 2003; Xue et al., 2014).



Figure 1. Time series of various air pollutants and meteorological parameters.

To simulate the concentration of particulate HCHO and its role in the heterogeneous formation of hydroxymethanesulfonate (HMS), the observation-based zero dimensional multiphase chemical box model was used, of which the gas-phase chemistry is described by the Regional Atmospheric Chemistry Mechanism, version 2 (RACM2), and the aqueous-phase part of the mechanism is represented by the Chemical Aqueous-Phase Radical Mechanism, version 3.0 (CAPRAM 3.0). The mass transfer processes between the gas and aqueous phases are also considered in the current model according to Schwartz (1986). The Henry's law constant of HCHO was updated with a value of $0.31 \times 10^8 \,\mathrm{M}\,\mathrm{atm}^{-1}$, as estimated by Mitsuishi et al. (2018). Sensitivity analysis was conducted to evaluate the uncertainties introduced by the Henry's law constant; details can be found in the Supplement. For major production and loss paths of HMS, dissolved HCHO reacts with sulfite and bisulfite to form HMS (Eqs. 3-4), which can be further oxidized by aqueous OH radicals (Eq. 5); details about the HMS mechanisms and the corresponding reaction kinetics are documented on the CAPRAM website (https://capram.tropos. de/, last access: 4 May 2023).

 $HCHO_{(aq)} + HSO_3^- = HOCH_2SO_3^-$ (3)

 $HCHO_{(aq)} + SO_3^{2-} + H_2O = HOCH_2SO_3^{-} + OH^{-}$ (4)

$$HOCH_2SO_3^- + OH \bullet = 2SO_4^{2-} + HCHO_{(aq)} + 3H^+$$
(5)

For detailed modeling steps of HMS, firstly, the observation data of gaseous NO, NO2, O3, SO2, CO, NMHCs (nonmethane hydrocarbons), HCHO, and 10 other carbonyls, particulate-phase NO_3^- , NH_4^+ , and Cl^- , along with meteorological parameters, were averaged or interpolated into a 1 h time resolution and classified into model recognized groups as model inputs, while the measured SO_4^{2-} was used as the initial conditions. Liquid water content (LWC) and aqueous H⁺ concentrations, calculated using the ISORROPIA-II model (Hong et al., 2022), were also used as model inputs. Then, model calculations were conducted using the commercial FACSIMILE software; the modeling period was from 26 February to 16 March 2022, and each day was regarded as an independent simulation case. The model was constrained every hour by the input observation data for integral calculation. For each case, the integration ran three times in series to steady the unconstrained species (e.g., radicals), which was initiated at 00:00 LT and had a step of 1 h and a duration of 24 h. Finally, the modeled HMS concentrations of the third run were outputted with a 1 h time resolution for further analysis.

3 Results and discussions

3.1 Overview of co-occurring O₃ and PM_{2.5} pollution

The time series of criteria air pollutants, O₃ precursors, and meteorological parameters from 17 February to 17 March 2022 are shown in Fig. 1. Two typical PM2.5 and O3 pollution episodes (EP1 – 26 February to 5 March; EP2 – 11 to 17 March) were observed, compared to the other periods (Pre-EP1 - 11 to 25 February; Pre-EP2 - 16 to 10 March) affected by rainfall. The mean concentrations of PM2.5 during EP1 and EP2 were 51.9 and $35.3 \,\mu g \,m^{-3}$, respectively, compared to 9.03 µg m⁻³ during Pre-EP1 (Table S1). The concentrations of other air pollutants, such as O₃, SO₂, NO₂, PM₁₀, OC, EC, BC, HCHO, and VOCs, showed a significant increasing trend during EP1 and EP2. During the monitoring periods, the concentrations of measured HCHO ranged from 0.68 and 3.59 ppbv (Table S1). According to our previous studies (Liu et al., 2023), the average levels of the measured HCHO in spring and autumn in Xiamen were 2.9 ± 0.3 and 3.2 ± 1.4 ppbv, respectively. In total, the HCHO level in Xiamen was lower than that in megacities (Table S2), such as Beijing (summer: 11.39 ± 5.58 ppbv), Hong Kong (summer: 8.07 ± 1.94 ppbv), and Guangzhou (summer: 6.69 ± 1.98 ppbv), while it was comparable to the coastal cities, including Shenzhen (spring: 3.4 ± 1.6 ppbv), Yantai (summer: 3.90 ± 1.12 ppbv), and Shanghai (summer: 3.31 ± 1.43 ppbv). In the coastal cities of southeastern China, halogenated VOCs are one of the most important VOC species, which originate from industrial emissions and solvent usage (Chen et al., 2022; Ji et al., 2022; Liu et al., 2022b). In this study, during the monitoring period, backward trajectories showed air mass transport from the northeast, which brought halogenated VOCs from the city of Quanzhou, an industrial city adjacent to Xiamen.

The maximum PM2.5 and O3 concentrations were approximately 100 and $200 \,\mu g \, m^{-3}$, respectively. The maximum daily 8 h average (MDA8) O3 concentrations were calculated according to the Ambient Air Quality Standard of China. Figure S3 shows the positive correlation between PM_{2.5} and MDA8 O₃ concentrations during the whole period. In Xiamen, a coastal city in southeast China, the annual mean concentrations of criteria air pollutants from 2015 to 2021 were significantly lower than in other Chinese cities (Fig. S4) (Li et al., 2022; Shao et al., 2022). Meanwhile, inter-annual averaged concentrations of O_3 and O_x in winter were shown in Fig. S4, suggesting an increase in atmospheric oxidation capacity (AOC) during the cold seasons. Therefore, these two typical PM_{2.5} and O₃ pollution episodes (EP1 and EP2) are worth exploring in terms of the formation mechanisms and synergistic effects of PM_{2.5} and O₃ in the coastal city.

3.2 Chemical compositions and sources of PM_{2.5}

As shown in Fig. 2, an overall increasing trend is clearly observed in both bulk PM2.5 and its major components during EP1 and EP2. Nitrate and organic matter (OM) remain the top two dominant PM2.5 components, followed by sulfate and ammonium. The mean concentrations of SO_4^{2-} , NO_3^{-} , and NH_4^+ during EP1 and EP2 were 7.07 and 5.87 µg m⁻³, 14.95 and 9.69 μ g m⁻³, and 6.77 and 4.46 μ g m⁻³, respectively (Table S1). The increase in EC indicates the contributions of local anthropogenic emission sources, such as vehicle exhausts (Fig. 2). The concentrations and percentages of OC and EC during different periods are illustrated in Fig. S5 and Table S1. The average OC and EC concentrations during EP1 and EP2 were 6.36 and 7.48 μ g m⁻³ and 1.23 and 1.29 μ g m⁻³, respectively, which were notably higher than those during Pre-EP1 and Pre-EP2. These results are consistent with the increase in primary emissions and secondary formation contributing to complex air pollution during the rapid urbanization and industrialization stages in China (Xiao et al., 2022; Jiang et al., 2022).

The PMF model was applied to conduct high-timeresolution source apportionment of PM2.5 based on online hourly measurement data (Hong et al., 2021; Chow et al., 2022). The factor profiles and the contributions of various sources to PM2.5 are shown in Figs. 2 and S6. Previous studies have indicated that construction and road dust is characterized by high loadings of Al, Si, Ca²⁺, Na⁺, Mg²⁺, and Zn (Rienda and Alves, 2021). In this study, the factor of dust (Factor 1) was identified by the high contributions of Si (Fig. S6). The PMF analysis revealed that the contribution of dust to PM2.5 ranged from 5.8 % to 8.9 % during EP1 and EP2, compared to 1.4 % during Pre-EP1 (Fig. 2). Factor 2, contributing to the high loading of metal elements (Mn, Zn, Fe, Pb, and As), was characterized by industrial emissions (Belis et al., 2019). The contributions from the industry during EP1 and EP2 remained constant. In factor 3, K⁺ was dominant, and it was identified as coming from combustion sources (Watson et al., 2001). Biomass burning could change the contribution of combustion to $PM_{2.5}$ at the monitoring site through long-range transport. During EP2, the influence of combustion sources (e.g., biomass burning) significantly decreased, due to reduced anthropogenic emissions and the arrival of clean air masses from the ocean (Figs. S7 and S8). Factor 4, with the highest proportion of Na⁺ and Mg²⁺ loadings, was associated with the influence of sea-salt aerosol (Polissar et al., 1998). The percentages of sea salt (8 %-10 %) during Pre-EP2 and EP2 were relatively high. Factor 5 exhibited high contributions of EC, OM, and Pb, which are general indicators of vehicle exhaust (Belis et al., 2019). During EP2, the contribution of traffic increased up to 35.3 %. Factor 6 was associated with secondary aerosol, characterized by high loadings of SO_4^{2-} , NO_3^- , and NH_4^+ . The increased contributions of secondary formation during EP1 and EP2 accounted for 47.8 % and 38.1 %, respectively.



Figure 2. Time series of PM_{2.5} chemical composition and sources apportionment by the PMF model. In the legends, OM refers to organic matter, calculated as $1.4 \times OC$; "Traces" includes elements other than Na, Cl, S, K, Al, Si, Ca, and Fe; "crustal" represents crustal materials, calculated as $1.89 \times Al + 2.14 \times Si + 1.4 \times Ca + 1.43 \times Fe$; and "salt" includes Na⁺ and Cl⁻.

3.3 Formation mechanism of PM_{2.5}

As shown in Fig. S9, SO_4^{2-} was correlated with NH_4^+ ($R^2 = 0.72-0.88$), and the line fit of NH_4^+ and SO_4^{2-} showed a slope of 1.78–2.67, suggesting the dominant form of (NH_4)₂SO₄. Similarly, NO_3^- was also correlated with NH_4^+ ($R^2 = 0.77-0.93$), indicating the presence of NH_4NO_3 . In addition, the ratio of NH_4^+ to the sum of NO_3^- and SO_4^{2-} was close to 1, indicating complete neutralization of sulfate and nitrate by ammonium (Fig. 8c). However, there was no significant difference in the existing form of SNA (sulfate, nitrate and ammonium) in PM_{2.5} during EP1 and EP2.

The variations of sulfur oxidation rate (SOR) and nitrogen oxidation rate (NOR) under different periods are shown in Table S1. It should be noted that SOR (0.38 ± 0.18) and NOR (0.32 ± 0.08) during EP1 were the highest, indicating a high oxidation rate of SO₂ and NO₂. According to RH, *T*, and UV (Table S1), noticeable differences in meteorological conditions were observed under different periods. In this study, LWC was positively correlated with SO²⁻₄, NO⁻₃, and NH⁺₄ (known as the secondary inorganic aerosol, SIA) (Fig. S10), suggesting the influence of the aqueous-phase process, including reactions with O₃, OH, H₂O₂, and organic peroxides (Gen et al., 2019; Wang et al., 2023).

Current studies have found that O_3 , H_2O_2 , OH, and transition metal ion (TMI)-catalyzed oxidation can trigger the secondary formation of SO_4^{2-} (Hong et al., 2021; Gen et al., 2019). However, the relative importance of these oxidants



Figure 3. Correlations between the concentrations of sulfate and HCHO (**a**), Fe (**b**), and Mn (**c**).

in enhancing the formation of SO_4^{2-} is still a topic of debate. As shown in Fig. 3b and c, a good correlation was found between SO_4^{2-} and Fe and Mn. The TMI-catalyzed oxidation contributed to the formation of SO_4^{2-} , which occurred in both cloud processes and during haze episodes (Li et al., 2020) because the Mn catalytic reaction rapidly occurred at the aerosol surface and could oxidize S(IV) through the production of intermediate Mn(III) (Wang et al., 2021).



Figure 4. Concentrations of SO₂, SO₄²⁻, and HCHO observed at different pollution stages. The simulated HCHO (aq) and HMS (aq) were also presented. The particulate sulfur molar percentage was calculated as $\left[n\left(SO_4^{2-}\right) + n(HMS)\right] / \left[n\left(SO_4^{2-}\right) + n(HMS) + n(SO_2)\right]$. In the box–whisker plots, the whiskers, boxes, and points indicate the 5th–95th, 25th–75th, and 50th percentiles and mean values.

Even at very low concentrations of Mn, the Mn catalytic reaction, consuming oxygen and SO₂, could produce sulfate. Under low aerosol pH conditions, the catalytic reaction of TMIs plays an important role in the oxidation of S(IV). In this study, a low aerosol pH ranging from 2 to 4 was observed, indicating the potential influence of TMI-catalyzed oxidation. In the future, it is vital to further evaluate the interaction of sulfate formation and Fe/Mn and to elucidate the main pathway of the S(IV) oxidation in the coastal areas. In addition, as an important intermediate product in atmospheric photochemical reactions, the formation and removal of HCHO are closely related to OH and HO₂ radicals, which directly affect atmospheric reactivity and oxidation ability (Wu et al., 2023; Zhang et al., 2021). In this study, the correlations ($R^2 = 0.415$) between HCHO and sulfate concentrations were also examined, as displayed in Fig. 3a. Recent studies have shown that HCHO can react with hydrogen peroxide (H_2O_2) to produce hydroxymethyl hydroperoxide, which rapidly oxidizes dissolved sulfur dioxide (SO_2, aq) to sulfate (Dovrou et al., 2022).

3.4 Effects of HCHO on HMS in PM_{2.5}

To calculate the particulate concentrations of HCHO and its contributions to the heterogeneous formation of HMS, we conducted the in situ multiphase chemical simulations by combining the zero dimensional multiphase chemical box model with RACM2 and CAPRAM 3.0. During EP1 and EP2, the concentrations of HCHO (aq) and HMS (aq), as well as the particulate sulfur molar percentage, increased with the rise of SO₂, SO₄²⁻, and HCHO concentrations (Fig. 4). Also, particulate sulfur molar percentage was higher during the pollution periods, which may be due to the high concentrations of gaseous precursors (SO₂, HCHO) and favorable aerosol properties (i.e., pH and aerosol water content) promoting the heterogeneous processes. Previous studies have modeled that 1 ppb of gas-phase HCHO could yield up to ~ 5 µg m⁻³ HMS and identified that HCHO emissions were one of the key factors controlling HMS formation during winter haze (Moch et al., 2018). All of the above signifies the important contributions of HCHO to the heterogeneous formation of HMS and particulate sulfur chemistry, especially during pollution periods.

However, the molar ratios of HMS to sulfate were very low, suggesting the limited contributions of HMS concentrations to inorganic sulfate concentrations. The potential role of HCHO in HMS formation in a coastal city of southeast China differed from that in megacities of China. Previous studies found that HCHO reacts with dissolved SO_2 (aq) to produce hydroxymethanesulfonate (HMS), which, upon oxidation with the hydroxyl radical (OH), forms sulfate (Ma et al., 2020; Moch et al., 2020). Ma et al. (2020) reported that heterogeneous formation of HMS accounted for 15 % of OM and resulted in 36 % overestimates of sulfate during winter haze in Beijing.

High concentrations of HMS were captured under high RH and moderately acidic pH conditions (Fig. 5). Previous studies have also indicated that high RH promotes rapid HMS



Figure 5. Evolution of HMS (aq) distribution with increasing RH, colored according to aerosol pH.

formation during winter haze, as the aerosol water content could provide numerous reaction interfaces for HMS formation (Ma et al., 2020). Meanwhile, atmospheric sulfur tended to distribute into the particle phase with increasing RH. Figure 5 shows that HMS formation is favored under pH conditions close to 4.0. Previous studies reported that high HMS concentrations were found under moderate-pH conditions, as low pH inhibits HMS formation, and high pH is unsuitable for its preservation (Ma et al., 2020; Campbell et al., 2022). Therefore, the combination of high precursor concentrations (SO₂ and HCHO), high RH, and moderately acidic pH enhanced the heterogeneous formation of HMS in the coastal city.

3.5 Effects of HCHO on O₃ formation

To investigate the effects of HCHO on O₃ formation during the co-occurring O_3 and $PM_{2.5}$ pollution period, the OBM was used to quantify the detailed O₃ production and loss pathways in both scenarios: input HCHO (IH) and noninput HCHO (NIH) (Figs. 6 and S10). The daytime production rates of $HO_2 + NO$ and $RO_2 + NO$ in the IH scenario were calculated to be 6.84 and 1.25 ppbv h^{-1} for EP1 and 9.91 and 2.17 ppbv h^{-1} for EP2, respectively. Meanwhile, the predominant O₃ loss reaction in this scenario was $OH + NO_2$, with rates of 2.26 ppbv h⁻¹ for EP1 and 3.17 ppbv h^{-1} for EP2, followed by O₃ photolysis with rates of 0.77 and 1.10 ppbv h^{-1} . In contrast, the daytime production rates of $HO_2 + NO$ and $RO_2 + NO$ in the NIH scenario were 4.03 and 0.85 ppbv h^{-1} for EP1 and 4.86 and 1.29 ppbv h^{-1} for EP2, respectively. These results indicate that disabling the HCHO mechanism reduced the production rates of $HO_2 + NO$ by 41 % for EP1 and 51 % for EP2. In addition, the average maximum net O₃ production rate with the IH scenario was $5.02 \text{ ppb } \text{h}^{-1}$ for EP1 and $7.93 \text{ ppb } \text{h}^{-1}$ for EP2, approximately 2 times higher than the values of 2.48



Figure 6. O₃ production and loss rates by the OBM with and without the HCHO mechanism.

and 3.14 ppb h⁻¹ with the NIH scenario. The results showed that the daytime net O₃ production rates decreased by 50 %– 60 % when the HCHO mechanism was disabled, probably due to the decrease in RO_x concentrations and radical propagation rates (Wu et al., 2023; Zhang et al., 2021).

The atmospheric oxidation capacity (AOC) is a critical factor in determining the production rate of secondary pollutants and atmospheric photochemical pollution (Jia et al., 2023; Qin et al., 2022). In this study, AOC is calculated as the sum of the oxidation rates of various primary pollutants (e.g., CO, NO_x , and VOCs) by major oxidants (i.e., OH, O₃, and NO₃). The model-simulated AOC, OH, HO₂, and RO₂ under different periods are shown in Figs. 7 and 8. The daily maximum AOC during EP1 and EP2 was 8.24×10^7 and 11.6×10^7 molec. cm⁻³ s⁻¹, respectively, which is higher than values (2.56×10^7) and 5.39×10^7 molec. cm⁻³ s⁻¹) in other periods. However, when the HCHO mechanism was disabled, the daily maximum AOC during different stages decreased significantly. In particular, HCHO played a very important role in AOC during the co-occurring PM_{2.5} and O₃ pollution periods. All these results are comparable to rural sites in Hong Kong (6.2×10^7) and Berlin $(1.4 \times 10^7 \text{ molec. cm}^{-3} \text{ s}^{-1})$ but lower than those observed in highly polluted cities,



Figure 7. Atmospheric oxidation capacity (AOC) calculated by the OBM with and without the HCHO mechanism.



Figure 8. OH, HO₂, and RO₂ concentrations modeled by the OBM with and without the HCHO mechanism.

such as Santiago $(3.2 \times 10^8 \text{ molec. cm}^{-3} \text{ s}^{-1})$ and Shanghai $(1.0 \times 10^8 \text{ molec. cm}^{-3} \text{ s}^{-1})$ (Li et al., 2018; Xue et al., 2016; Liu et al., 2022a). These studies have reported that the variations in AOC are related to precursor concentrations/types and photochemical conditions.

In addition, the maximum daily concentrations of OH, HO_2 , and RO_2 exhibited a similar pattern to that of AOC in both the IH and NIH scenarios (Fig. 8). Therefore, the O₃ production rate during EP1 and EP2 was consistent with the maximum daily values of AOC, OH, HO_2 , and RO_2 . The dif-

ferences in RO_x levels between the IH and NIH model scenarios were also calculated (Fig. S11). In this study, disabling the HCHO mechanism led to decreased RO_x concentrations, affecting the O₃ formation. These results highlight the significance of HCHO in the photochemical reactions occurring in the coastal city during the co-occurring O₃ and PM_{2.5} pollution events.

4 Conclusions

A wintertime co-occurring O₃ and PM_{2.5} pollution event was selected to investigate the synergistic effects between $PM_{2.5}$ and O_3 in a coastal city in southeast China. The results demonstrated a positive correlation between PM_{2.5} and MDA8 O₃ concentrations during the whole period, indicating the enhancement of atmospheric oxidation capacity (AOC) during cold seasons. The result of positive matrix factorization (PMF) analysis suggested that the contribution of secondary formation to PM2.5 increased during the pollution events, implying that the elevated AOC promoted the oxidation of SO_2 , NO_x , and VOCs, leading to the formation of secondary inorganic and organic components. We also found significant correlations ($R^2 = 0.415 - 0.477$) between HCHO, Fe, Mn, and sulfate concentrations, suggesting the influence of catalyzed oxidation in the coastal city. Through OBM analysis, we demonstrated that high concentrations of precursors (SO₂ and HCHO), high RH, and moderately acidic pH conditions enhanced the heterogeneous formation of hydroxymethanesulfonate (HMS). Meanwhile, we verified that the input HCHO mechanism increased the concentrations of RO_x and net O_3 production rates. Moreover, the production rates of $HO_2 + NO$ and $RO_2 + NO$ were enhanced, indicating that HCHO affected O₃ formation by controlling the efficiencies of radical propagation. This study highlighted the influence of the HCHO mechanism on photochemical reactions and the formation of HMS in PM_{2.5} in coastal cities and was beneficial for improving air quality and protecting public health.

Data availability. The dataset related to this work can be accessed via https://doi.org/10.5281/zenodo.7799302 (Hong, 2023). Details are also available upon request from the corresponding author (ywhong@iue.ac.cn).

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